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The following Figure shows the size ranges of particles that play a role in the atmosphere.



The electrical conductivity of the air is generally dominated by small ions. However, when the concentrations of large ions and uncharged aerosols are large, as they are in cities, the concentration of small ions tends to be low due to their capture by both large ions and uncharged aerosols.

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Consequently, the electrical conductivity of air is a minimum when the concentration of large ions is a maximum.

The observed decrease of at least 20% in the electrical conductivity of the air over the North Atlantic Ocean during the 20^{th} century is attributed to a doubling in the concentration of particles with diameters between 0.02 and 0.2 μ m, probably due to pollution.



Sources

a. Biological

Solid and liquid particles are released into the atmosphere from plants and animals. These emissions, which include seeds, pollen, spores, and fragments of animals and plants, are usually 1-250 μ m in diameter. Bacteria, algae, protozoa, fungi, and viruses are generally <1 μ m in diameter.

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Microorganisms live on skin: when you change your clothes, you can propel $\sim 10^4$ bacteria per minute into the air, with diameters from 1–5 μ m.

The oceans are an important source of biogenic aerosols. They are injected into the atmosphere by the bursting of air bubbles and by sea foam.



Production of film droplets and jet drops when an air bubble bursts at the surface of water. The time between (a) and (d) is ~ 2 ms.

* * *



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The number distribution of particles from forest fires peak at $\sim 0.1 \,\mu\text{m}$ diameter, which makes them efficient cloud condensation nuclei. Some biogenic particles (e.g., bacteria from vegetation) may nucleate ice in clouds.

b. Solid Earth

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Volcanoes inject gases and particles into the atmosphere. The large particles have short residence times, but the small particles can be transported globally, particularly if they are blasted to high altitudes.

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It is projected that by 2040, anthropogenic sources of particles could be comparable to those from natural processes.

In situ condensation of gases (i.e., g-to-p conversion) is important in the atmosphere. Gases may condense onto existing particles, thereby increasing the mass (but not the number) of particles, or gases may condense to form new particles.

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Summary: Table 5.3 in *Wallace & Hobbs* summarizes the magnitudes of the principal sources of direct emission of particles into the atmosphere and in situ sources.

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In marine air the main contributors to the mass of inorganic aerosols are the ions Na⁺, Cl⁻, Mg²⁺, SO₄²⁺, K⁺ and Ca²⁺. Such particles are particularly effective in scattering light (reducing visibility) and as cloud condensation nuclei.

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Nitrate (NO_3^-) occurs in larger sized particles than sulfate in marine air. Since seawater contains negligible nitrate, the nitrate in these particles must derive from the condensation of gaseous HNO₃ by g-to-p conversion. Many of the particles in urban smog are by-products of photochemical reactions involving hydrocarbons and nitrogen oxides, which derive from combustion.
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Elemental carbon (commonly referred to as "soot") is a common component of organic aerosols in the atmosphere, and is a strong absorber of solar radiation.

For example, in polluted air masses from India, elemental carbon contributes about 10% to the mass of submicrometer sized particles.

Transport

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Thus, Saharan dust is transported to the Americas, and dust from the Gobi Desert can reach the west coast of North America.

Sulpher dioxide (SO_2) emitted from power plants in the United Kingdom can be deposited as sulfate far inland in continental Europe.

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Improvements in visibility that frequently follow periods of precipitation are due, in large part, to the removal (i.e., scavenging) of particles by precipitation. It is estimated that, on the global scale, precipitation processes account for about 80-90% of the mass of particles removed from the atmosphere.

The terminal fall speeds of particles >1 μ m diameter are sufficiently large that gravitational settling (i.e., dry deposition) is important as a removal process. The terminal fall speeds of particles >1 μ m diameter are sufficiently large that gravitational settling (i.e., dry deposition) is important as a removal process.

For example, the fall speeds of particles 1 and 10 μ m in diameter are ~ 3 × 10⁻⁵ and ~ 3 × 10⁻³ m s⁻¹, respectively.

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For example, the fall speeds of particles 1 and 10 μ m in diameter are ~ 3 × 10⁻⁵ and ~ 3 × 10⁻³ m s⁻¹, respectively.

Some 10-20% of the mass of particles removed from the atmosphere is by dry fallout.

Concentrations.

One of the oldest and most convenient techniques (which in various forms is still in widespread use) for measuring the concentrations of particles in the air is the Aitken nucleus counter. For details, see *Wallace & Hobbs*.

The concentration of particles measured with this instrument is referred to as the Condensation Nucleus (CN) count.

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Condensation nucleus counts near the Earth's surface vary widely from one location to another, and they can also fluctuate by more than an order of magnitude with time at any one site.

Generally, they range from average values of $\sim 10^3 \, \text{cm}^{-3}$ over the oceans, to $\sim 10^4 \, \text{cm}^{-3}$ over rural land areas, to $\sim 10^5 \, \text{cm}^{-3}$ or greater in urban polluted air.

Size Distributions

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The measurements are plotted in the form of a number distribution in which the ordinate $dN/d(\log N)$ and the abscissa (D) are plotted on logarithmic scales, where dN is the number concentration of particles with diameters between D and D + dD.



Number distributions of tropospheric particles in continental (red), marine (blue), and urban polluted (black) air.

Conclusions from results shown in Figure.

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- Those portions of the number distribution curves that are straight lines in the Figure can be represented by an expression of the form

$$\log \frac{dN}{d(\log D)} = \mathbf{const} - \beta \log D$$

or, taking antilogs,

$$\frac{dN}{d(\log D)} = CD^{-\beta}$$

• The number distributions of particles shown confirm CN measurements, which indicate that the total concentrations of particles are, on average, greatest in urban polluted air and least in marine air.

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- The concentrations of particles with diameters >2 μ m (giant particles) are, on average, rather similar in continental, marine, and urban polluted air.

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Dispersion in the vertical is particularly restricted under stable conditions.

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Ideal (or complete) combustion of a hydrocarbon fuel yields only CO_2 and H_2O . However, for a given quantity of fuel, a precise amount of oxygen is required for complete combustion, and this ideal combination of fuel and oxygen is rarely achieved.

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There are also lower temperature sources of air pollutants; for example, leakages of hydrocarbons from natural gas lines, organics from the evaporation of solvents, and nitrogen gases from fertilizers. Anthropogenic activities also emit large numbers of particles into the atmosphere, both directly and through g-to-p conversion. Anthropogenic activities also emit large numbers of particles into the atmosphere, both directly and through g-to-p conversion.

For particles $>5 \,\mu\text{m}$ diameter, human activities worldwide are estimated to produce $\sim 15\%$ of natural emissions, with industrial processes, fuel combustion, and g-to-p conversion accounting for $\sim 80\%$ of the anthropogenic emissions. Anthropogenic activities also emit large numbers of particles into the atmosphere, both directly and through g-to-p conversion.

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Sulfur dioxide gas absorbs into the fog droplets where it is oxidized to form sulfuric acid.

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In the Great Smog, as it was called, SO₂ reached peak mixing ratios of about 0.7 ppmv (compared to typical annual mean mixing ratios of about 0.1 ppmv in polluted cities with large coal usage), and the peak particle concentrations were 1.7 mg m⁻³. After the Great Smog, laws were passed in Britain and elsewhere banning the use of coal on open fires for domestic heating and the emissions of black smoke, and requiring industries to switch to cleaner-burning fuels. After the Great Smog, laws were passed in Britain and elsewhere banning the use of coal on open fires for domestic heating and the emissions of black smoke, and requiring industries to switch to cleaner-burning fuels.

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Thank you, Mary Harney!

Nevertheless, pollution is still prevalent in significant quantities in many cities in Europe and the United States. Also, many large cities, particularly in developing countries (e.g., China, India), still suffer from London-type smogs due to the burning of coal and wood and to the lack of strict air pollution controls.

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These smogs are characterized by high concentrations of a large variety of pollutants, such as nitrogen oxides, O_3 , CO, hydrocarbons, aldehydes (and other materials that are eye irritants), and sometimes sulfuric acid.

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These smogs are characterized by high concentrations of a large variety of pollutants, such as nitrogen oxides, O_3 , CO, hydrocarbons, aldehydes (and other materials that are eye irritants), and sometimes sulfuric acid.

The chemical reactions that lead to photochemical smog are complex, and still not completely understood. For an outline, see *Wallace & Hobbs*.

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Europe, Russia, the northeastern United States, India, and large areas of southeastern Asia, are regularly covered by enormous palls of polluted air that reduce visibility significantly, produce acid deposition, soil and erode buildings and other materials, and have deleterious effects on human health, animals and plants.

The fact that pollutants can be transported over large distances is well illustrated by air pollution episodes in the Arctic, known as arctic haze, which can be as severe as those in cities. The pollutants originate from fossil-fuel combustion, smelting, and other industrial processes in northern Europe and Russia. The pollutants are transported to the Arctic by synopticscale airflow patterns, primarily from December to April. Since the arctic atmosphere is generally stably stratified during this time of the year, vertical mixing is limited; also, precipitation is rare so that wet removal processes are weak. The pollutants are transported to the Arctic by synopticscale airflow patterns, primarily from December to April. Since the arctic atmosphere is generally stably stratified during this time of the year, vertical mixing is limited; also, precipitation is rare so that wet removal processes are weak.

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Glacial records show that air pollution in the Arctic has increased markedly since the 1950s, paralleling the increases in SO_2 and NO_x emissions in Europe.

Ice cores from Greenland show unusually high lead concentrations from ~ 500 B.C. to 300 A.D. This is attributed to Greek and Roman lead and silver mining and smelting activities, which apparently polluted large regions of the Northern Hemisphere. Ice cores from Greenland show unusually high lead concentrations from ~ 500 B.C. to 300 A.D. This is attributed to Greek and Roman lead and silver mining and smelting activities, which apparently polluted large regions of the Northern Hemisphere.

Anthropogenic influences are now apparent on a global scale, as illustrated by the world-wide increase in CO_2 concentrations since the Industrial Revolution (following Figure). Other trace gases (e.g., CH_4 and N_2O) also show increasing concentrations worldwide over the past 150 years or so.



Changes in the concentrations of (a) CO_2 , (b) CH_4 , and (c) N_2O over 300 years deduced from analyses of ice cores.

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The reservoirs of chemical species in the Earth system are the solid Earth, the hydrosphere, the cryosphere, the biosphere and the atmosphere. Chemical species can be transferred between these reservoirs.

Since, under steady-state conditions, a chemical species cannot accumulate indefinitely in any of the reservoirs, there must be continual cycling of species through the various reservoirs. This is termed biogeochemical cycling.

Here we will consider the tropospheric portions of the biogeochemical cycles of nitrogen and sulfur. We will be concerned with relatively rapid interchanges involving the atmosphere and other reservoirs (generally the oceans and the biosphere). Many aspects of global chemical cycles are not well understood, therefore, in many cases, the magnitudes of the sources and sinks given are estimates. Many aspects of global chemical cycles are not well understood, therefore, in many cases, the magnitudes of the sources and sinks given are estimates.

Also, the magnitudes of the various emission fluxes are not necessarily equivalent to the importance of the species, since atmospheric residence times must also be taken into account.

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The other nitrogen species in the atmosphere are therefore present in very low concentrations, but they are of crucial importance in atmospheric chemistry.

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Nitric oxide and NO_2 play important roles in both tropospheric and stratospheric chemistry.
The principal sources and sinks of nitrogen-containing species in the atmosphere are shown in the following Figure. The main atmospheric sources of nitrogen-containing species are biogenic emissions from the Earth and the oceans (NH₃, N₂O, and NO_x), decomposition of proteins and urea from animals (NH₃), biomass burning and fossil fuel consumption (NO_x, NH₃, and N₂), and lightning (NO_x). The principal sources and sinks of nitrogen-containing species in the atmosphere are shown in the following Figure. The main atmospheric sources of nitrogen-containing species are biogenic emissions from the Earth and the oceans (NH₃, N₂O, and NO_x), decomposition of proteins and urea from animals (NH₃), biomass burning and fossil fuel consumption (NO_x, NH₃, and N₂), and lightning (NO_x).

The main sinks are wet removal by precipitation (NH₃ and NO_x as NO₃⁻), dry deposition (NO_x and NH₃), and the chemical breakdown of N₂O in the stratosphere.

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Since anthropogenic sources of NH_3 , N_2O , and NO_x (from fossil fuel consumption, biomass burning, and agricultural nitrate fertilization) are appreciable, they may be causing significant perturbations in the budgets of these species in the atmosphere.



Principal sources and sinks of nitrogen-containing gases in the atmosphere.

The most important reduced sulfur gases in the air are H_2S , DMS, COS, and CS₂. Their main natural sources are biogenic reactions in soils, marshland and plants.

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Biogenic reactions in the oceans, due primarily to phytoplankton, are sources of DMS, COS and CS₂. Estimates of the fluxes of these natural emissions of sulfur gases, and their transformations to SO_2 and SO_4^{2-} are given in the figure below. DMS dominates the emissions of sulfur gases from the oceans.

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The global sulfur budget is significantly affected by anthropogenic emissions. These are almost entirely in the form of SO_2 , and 90% are from the Northern Hemisphere. The main sources are the burning of coal and the smelting of sulfide ores.

The main mechanisms for removing sulfur from the atmosphere are wet and dry deposition. The main mechanisms for removing sulfur from the atmosphere are wet and dry deposition.

For example, of the 80 Tg(S) per year of SO₂ that are oxidized to SO₄²⁻, about 70 Tg(S) per year occurs in clouds, which is subsequently wet deposited.



Principal sources and sinks of sulpher-containing gases in the atmosphere.

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where M is the amount of the constituent in the atmosphere and F the rate of its removal from the atmosphere.

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An analogy is a tank of water, which can represent the atmosphere. Assume water is pumped into and removed from the tank at rate F. The mean time spent by each small element of water is M/F, where M is the volume of the tank.



Spatial and temporal scales of variability for some atmospheric constituents. The temporal scale is represented by residence time. If a chemical species has a very short residence time in the atmosphere, significant variations in the concentration of the species will generally occur over very short spatial scales. If a chemical species has a very short residence time in the atmosphere, significant variations in the concentration of the species will generally occur over very short spatial scales.

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The residence times of particles in the atmosphere as a function of their size are shown below.



Above: Schematic curves of particle surface area distributions for urban polluted air (black line), continental air (red line) and marine air (blue line).

Below: principal sources and sinks of atmospheric particles and estimates of their mean residence times in the troposphere.