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Nudging is still used for mesoscale analysis, for example, using radar data.

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Large scales are analysed first, then smaller scales.

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The variance of pressure,  $\sigma_p^2$ ,

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The standard deviations,  $\sigma_p$  and  $\sigma_T$  are the square roots of the variances. They measure the root mean square deviation from the mean.

3

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If p tends to be greater than its mean value when T is less than its mean, and vice-versa, then p and T are negatively correlated and  $\rho_{pT} < 0$ .

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The mean is  $\bar{x}$  and the standard deviation  $\sigma$ . The constant term is included so that  $\int_{-\infty}^{+\infty} P(x) dx = 1$ .

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Many variables have distributions which can be accurately modelled by the normal distribution.

## Back to the SCM Analysis



Figure 5.1.1: Schematic of grid points (circles), irregularly distributed observations (squares), and a radius of influence around a grid point *i* marked with a black circle. In 4DDA, the grid-point analysis is a combination of the forecast at the grid point (first guess) and the observational increments (observation minus first guess) computed at the observational points k. In certain analysis schemes, like SCM, only observations within the radius

of influence, indicated by a circle, affect the analysis at the black grid point.

$$f_{i}^{n+1} = f_{i}^{n} + \begin{bmatrix} \sum_{k=1}^{K_{i}^{n}} w_{ik}^{n} (f_{k}^{O} - f_{k}^{n}) \\ \frac{k=1}{\sum_{k=1}^{K_{i}^{n}} w_{ik}^{n} + \varepsilon^{2}} \end{bmatrix}$$

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The notation is as follows:

•  $f_i^n$  is the *n*th iteration estimation at the grid point *i* 

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- $f_i^n$  is the *n*th iteration estimation at the grid point *i*
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- $K_i^n$  is the number of observations within a distance  $R_n$  of the grid point i
- $\varepsilon^2$  is an estimate of the ratio of the observation error variance to the background error variance.

To calculate the difference  $f_k^O - f_k^n$ , we have to interpolate the background field to the observation point.



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We interpolate first in the x-direction, and then interpolate the results in the y-direction.

Usually, bi-quadratic or bi-cubic interpolation is used.

Cressman (1959) defined the weights  $w_{ik}^n$  in the SCM as

$$w_{ik}^{n} = \frac{R_{n}^{2} - r_{ik}^{2}}{R_{n}^{2} + r_{ik}^{2}} \quad \text{for} \quad r_{ik}^{2} \le R_{n}^{2}$$
$$w_{ik}^{n} = 0 \quad \text{for} \quad r_{ik}^{2} > R_{n}^{2}$$

where  $r_{ik}^2$  is the square of the distance between an observation point  $\mathbf{r}_k$  and a grid point at  $\mathbf{r}_i$ .

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**Exercise:** Consider a single observation at a gridpoint. If the first guess is a uniform field, what does the SCM analysis look like? (Assume  $\varepsilon = 0$ ). Cressman (1959) defined the weights  $w_{ik}^n$  in the SCM as

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**Exercise:** Consider a single observation at a gridpoint. If the first guess is a uniform field, what does the SCM analysis look like? (Assume  $\varepsilon = 0$ ).

**Exercise:** Draw a series of weight curves using MatLab.



Figure 5.1.1: Schematic of grid points (circles), irregularly distributed observations (squares), and a radius of influence around a grid point *i* marked with a black circle. In 4DDA, the grid-point analysis is a combination of the forecast at the grid point (first guess) and the observational increments (observation minus first guess) computed at the observational points k. In certain analysis schemes, like SCM, only observations within the radius

of influence, indicated by a circle, affect the analysis at the black grid point.

There are eight observations within a distance R from grid point i. The radius of influence  $R_n$  may vary with the iteration.

For example, the Swedish operational system (in 1980s) used

- $R_1 = 1500$  km,  $R_2 = 900$  km for upper air analyses
- $R_1 = 1500$  km,  $R_2 = 1200$  km,  $R_3 = 750$  km,  $R_4 = 300$  km for the surface pressure analysis

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The field reflects the large scales after the first iteration, and converges towards the smaller scales during later iterations.

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**Exercise:** Consider the effect of a single observation, located at a greid point (a) for  $\varepsilon^2 = 0$ , (b) for  $\varepsilon^2 = 0.5$ .

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For example, for a primitive equation model, the zonal velocity forecast equation is written as

$$\frac{\partial u}{\partial t} = -\mathbf{v} \cdot \nabla u + fv - \frac{\partial \phi}{\partial x} + \left[\frac{u_{obs} - u}{\tau_u}\right]$$

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The relaxation time scale,  $\tau$ , is chosen based on empirical considerations.

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Recall how to solve this o.d.e. using an integrating factor. Multiply by  $\exp\left(\int dt/\tau\right) = \exp(t/\tau)$ :

$$\frac{d}{dt} \left[ \exp\left(\frac{t}{\tau}\right) u \right] = \exp\left(\frac{t}{\tau}\right) \frac{u_{obs}}{\tau}$$

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So, u approaches  $u_{obs}$  exponentially, with time-scale  $\tau$ .

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Kaas *et al.* (1999) performed an interesting experiment, nudging a model towards a 15-y reanalysis from the ECMWF.

By averaging the mean forcing due to nudging, they empirically determined corrections to reduce model deficiencies. **Question:** How is the observational information *spread out in space* in the nudging analysis scheme?

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Some groups use it for assimilating small-scale observations (e.g., radar observations).

It is especially useful when there are no available statistics to perform a statistical interpolation.

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**Exercise:** Assuming that the nudging term is comparable in size to the Coriolis term, estimate the relaxation time  $\tau$ .