

Numerical Weather Prediction

Prof Peter Lynch

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School of Mathematical Sciences
University College Dublin
Second Semester, 2005–2006.*

Text for the Course

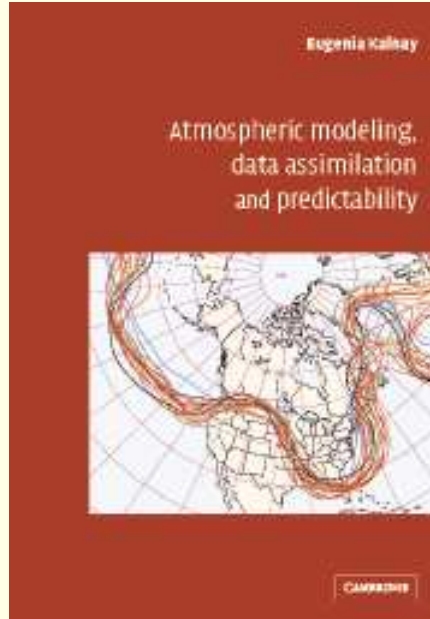
The lectures will be based closely on the text

Atmospheric Modeling, Data Assimilation and Predictability

by

Eugenia Kalnay

published by Cambridge University Press (2002).



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- This approach is called **data assimilation**

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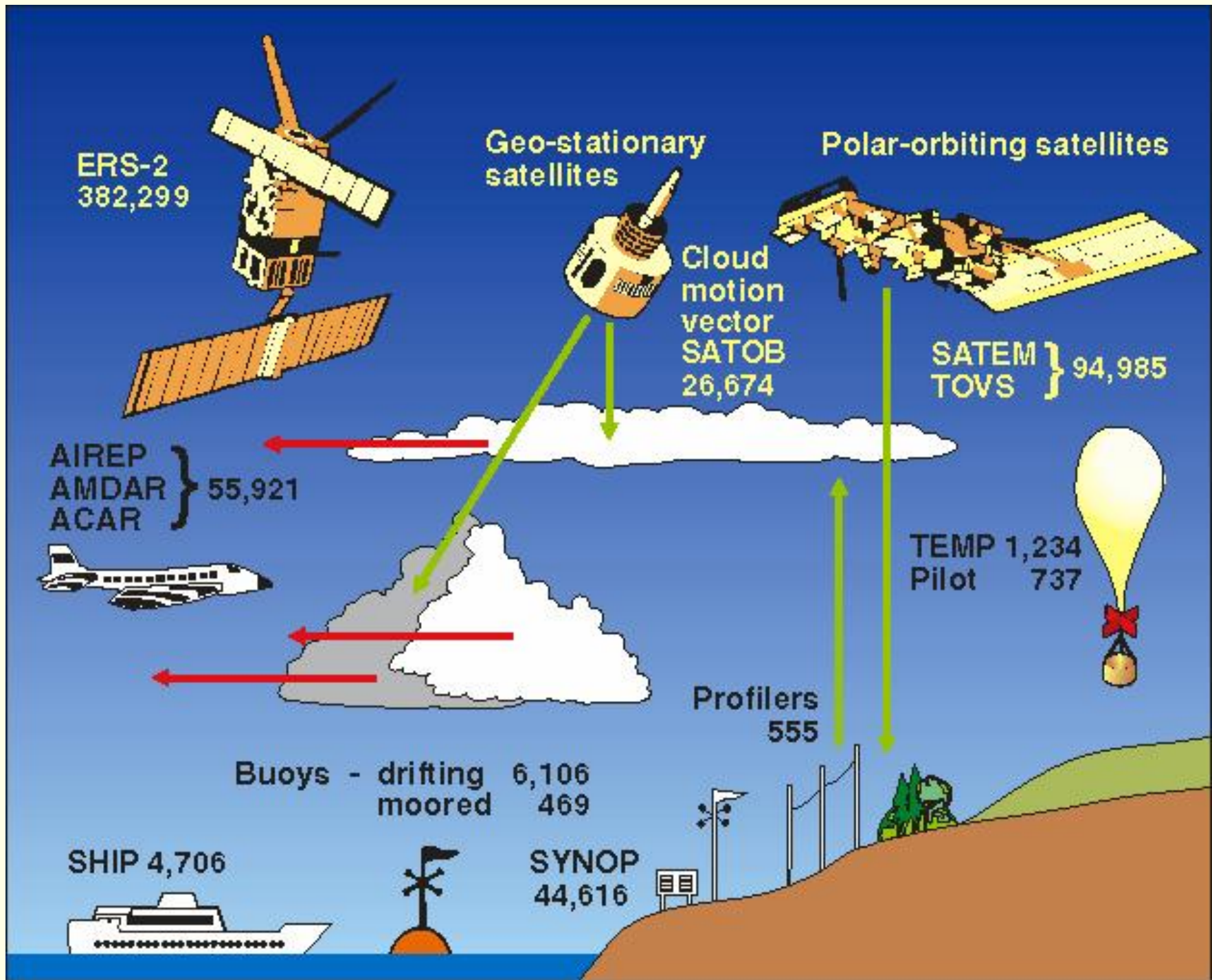
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The first objective analysis systems were developed (independently) in Sweden and in USA in the 1950s.

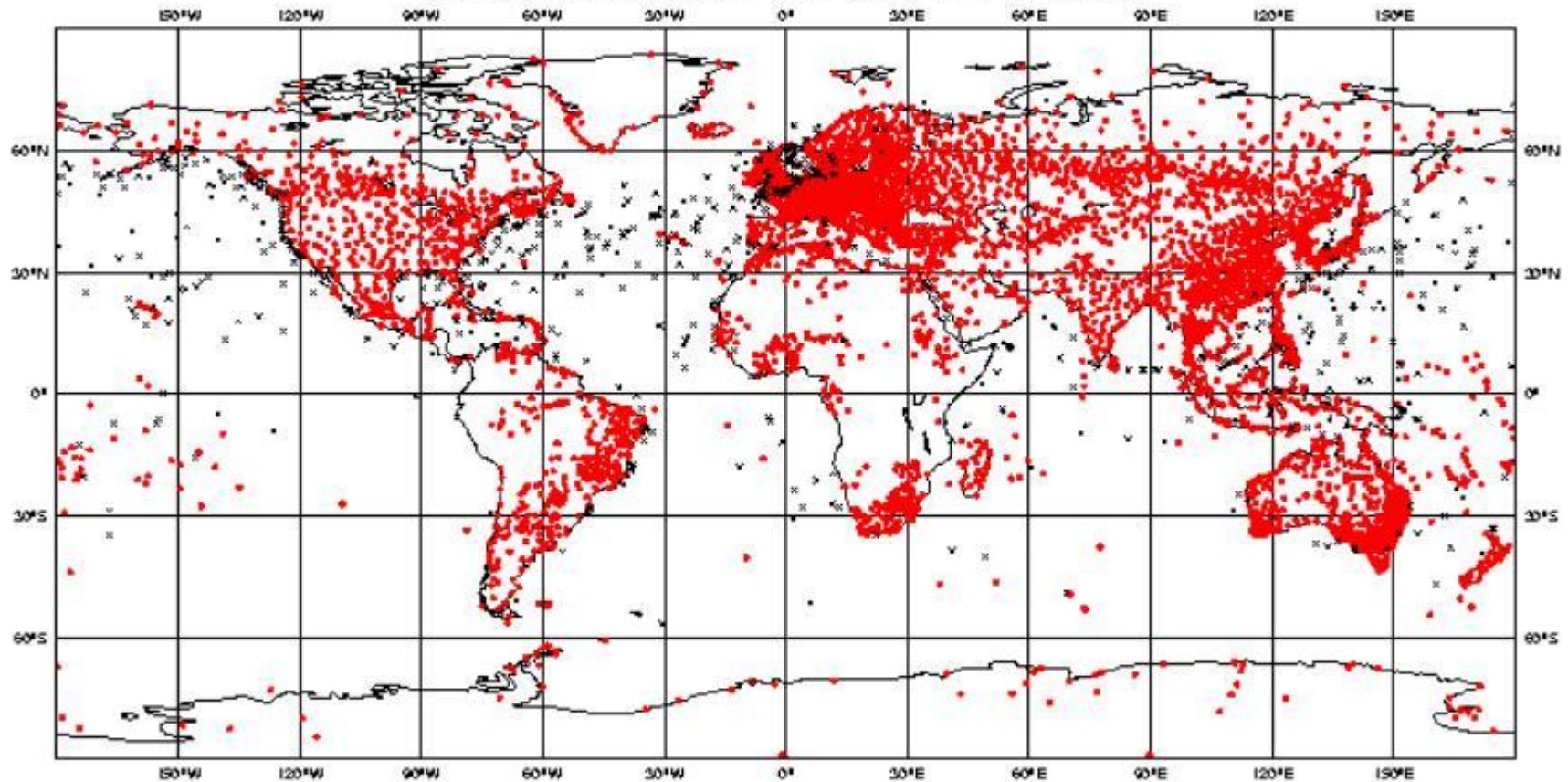


ECMWF Data Coverage - SYNOP/SHIP

28/FEB/1999; 00 UTC

Total number of obs = 12688

- 11417 SYNOP
- 1271 SHIP

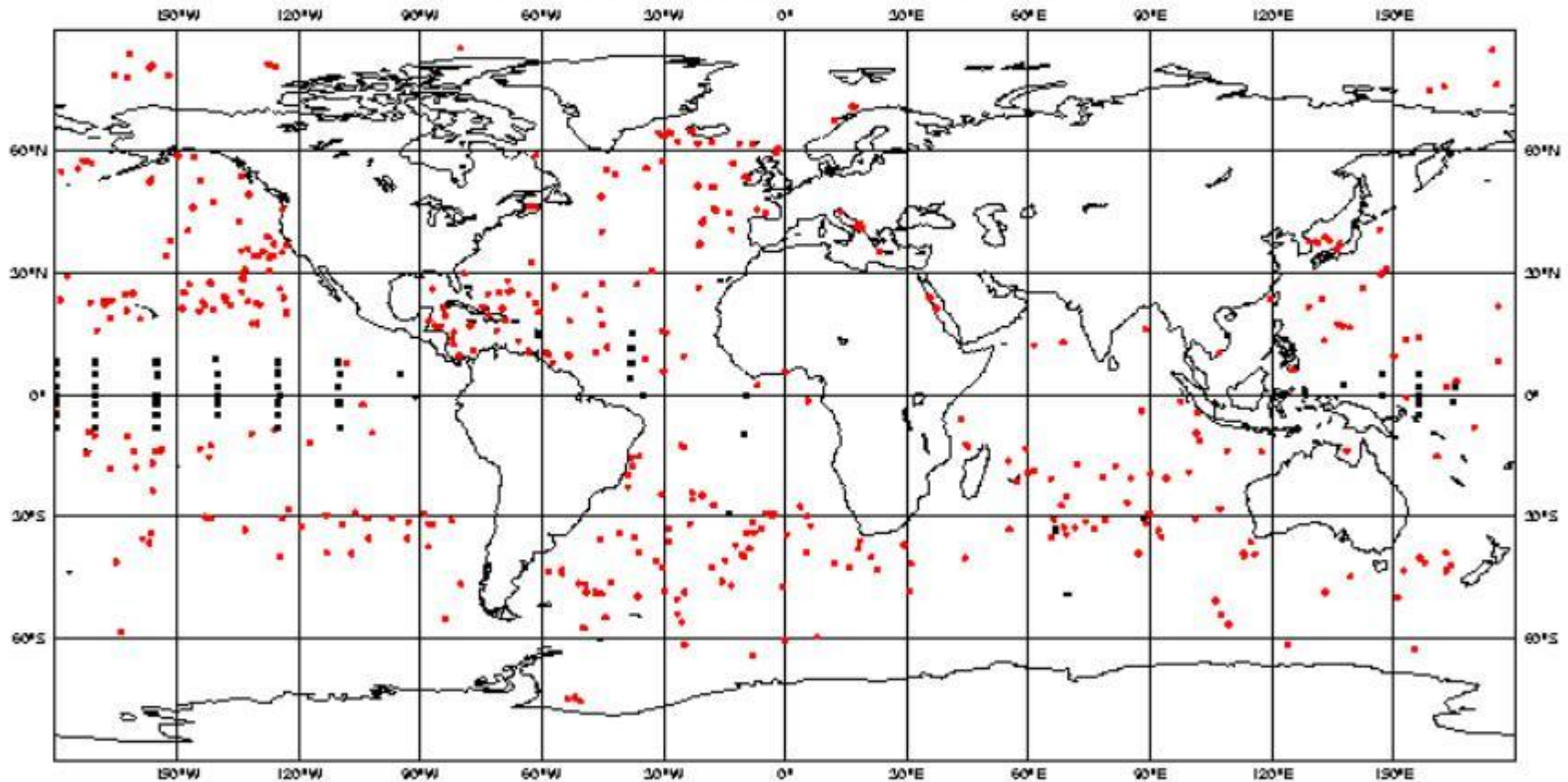


ECMWF Data Coverage - BUOY

28/FEB/1999; 00 UTC

Total number of obs = 1568

- 1488 DRIFTER
- 100 MOORED

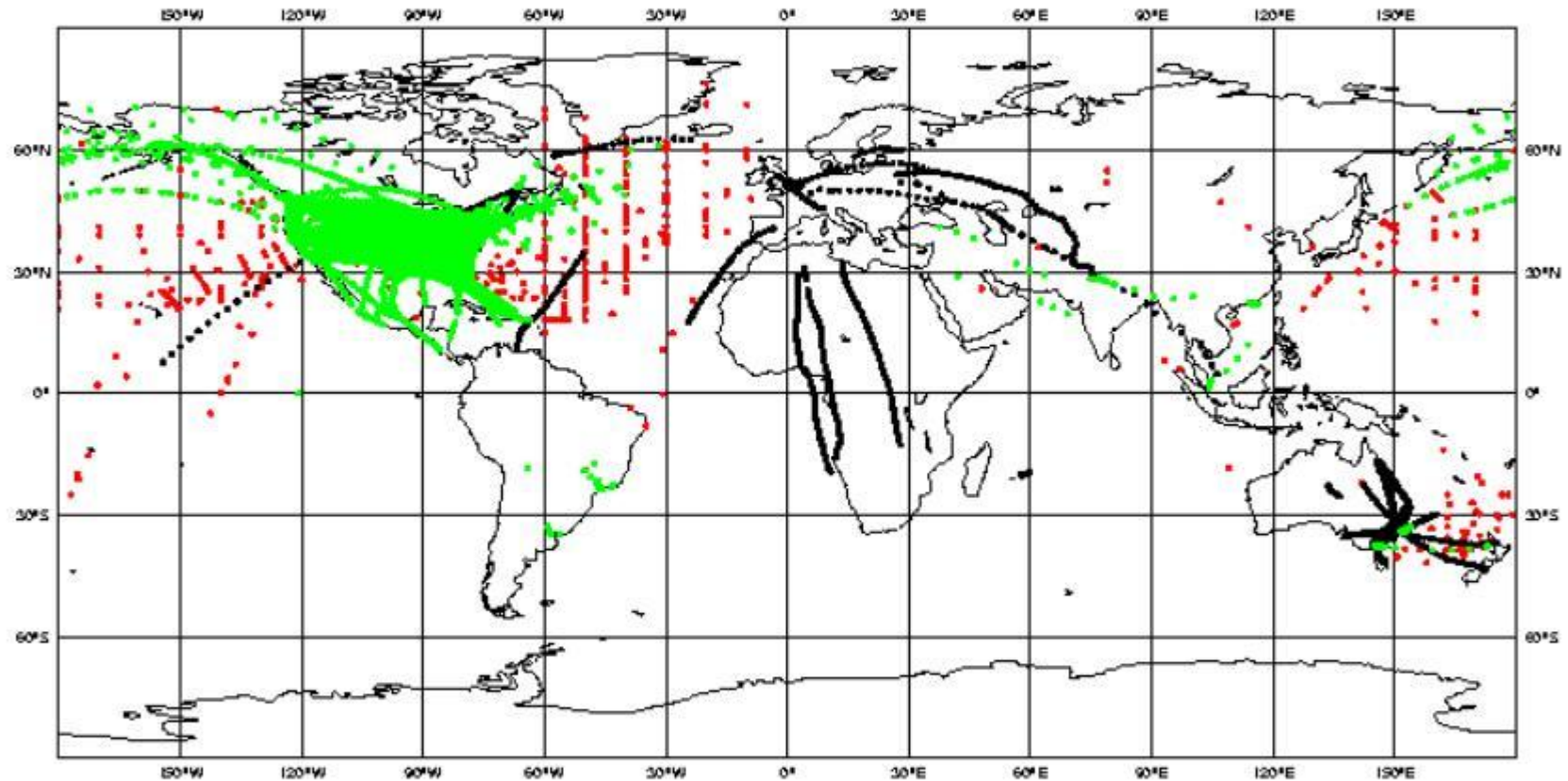


ECMWF Data Coverage - AIRCRAFT

28/FEB/1999; 00 UTC

Total number of obs = 18964

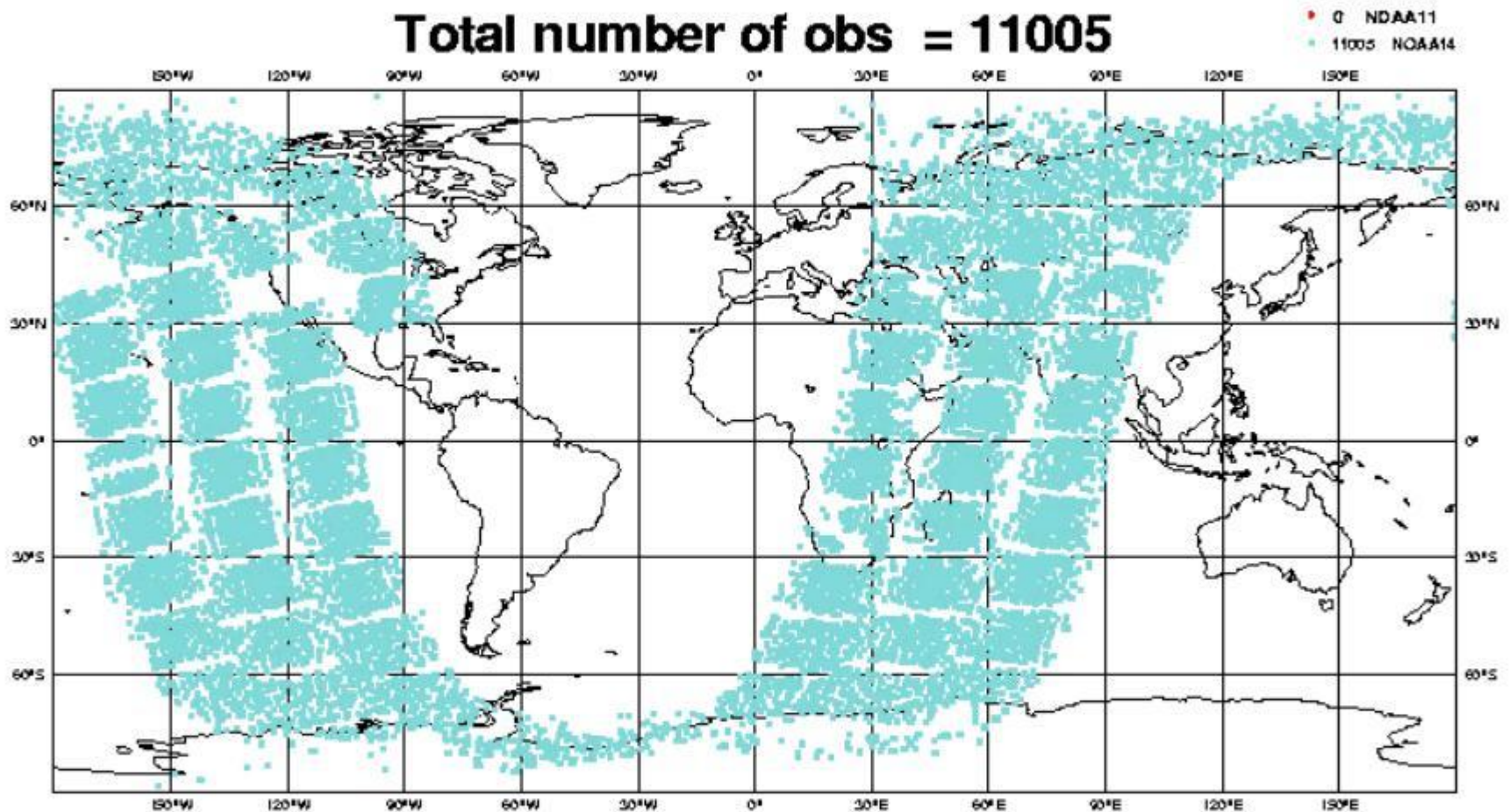
- 5726 AIREP
- 2346 AMDAR
- 10892 ACARS



ECMWF Data Coverage - TOVS (120km)

28/FEB/1999; 00 UTC

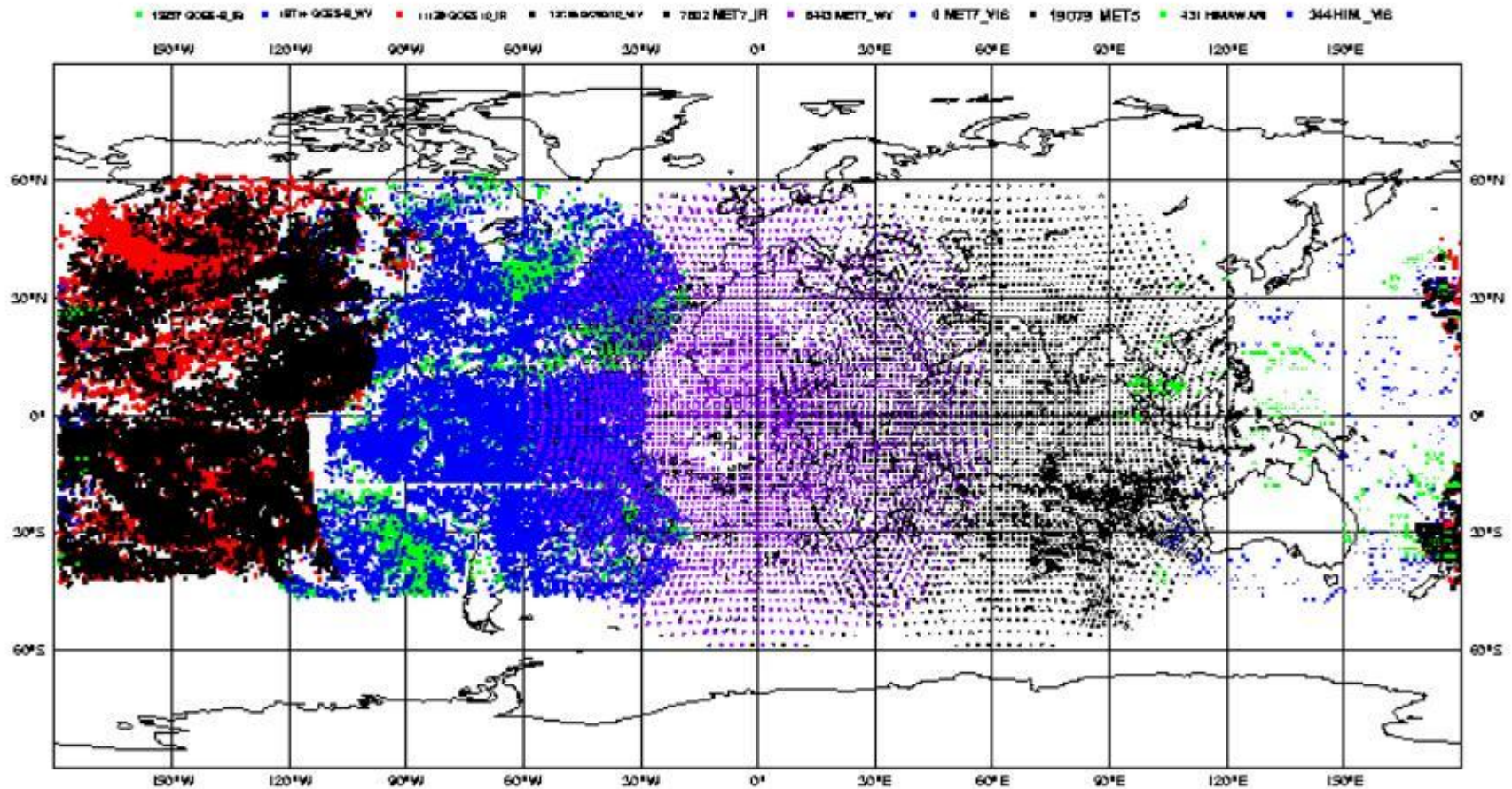
Total number of obs = 11005



ECMWF Data Coverage - SATOB

28/FEB/1999; 00 UTC

Total number of obs = 91405

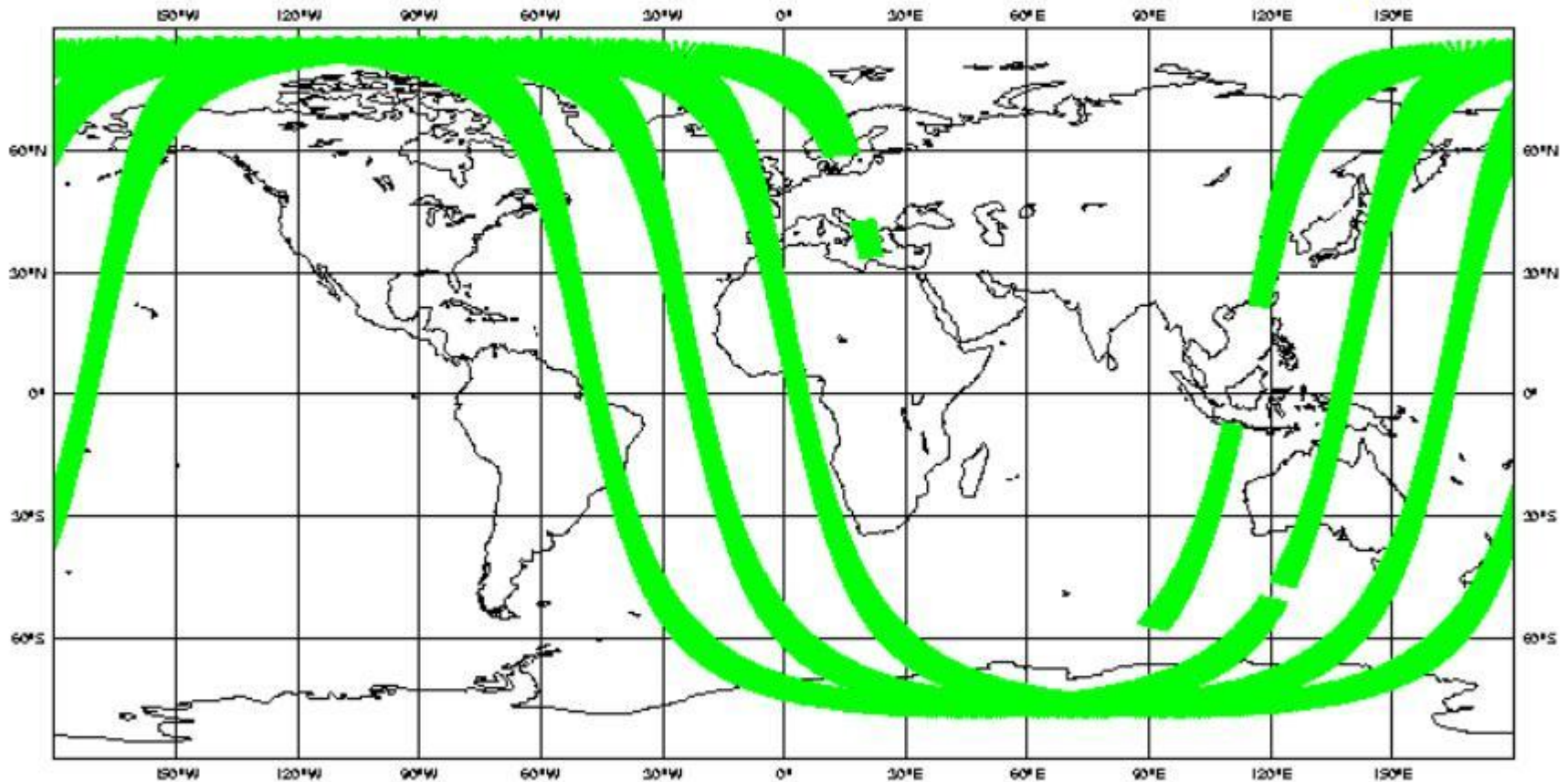


ECMWF Data Coverage - ERS-2

28/FEB/1999; 00 UTC

Total number of obs = 107939

- 107939 AUTO.
- 0 MET.
- 0 RANK_1
- 0 NS0=3
- 0 NS0=2
- 0 NS0=1
- 0 NS0=0



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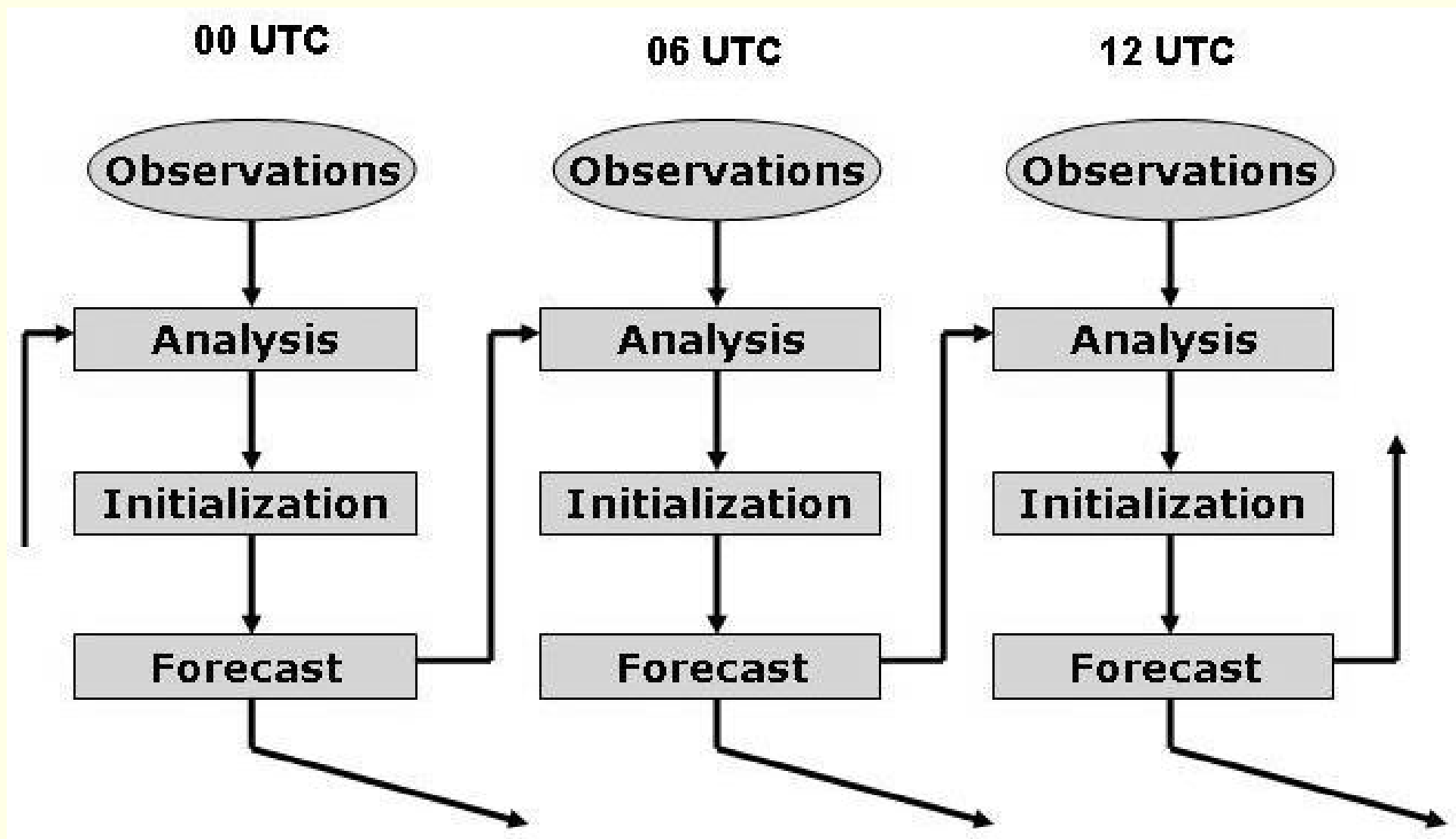
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Present-day operational systems typically use a 6-h cycle performed four times a day.



Typical 6-hour analysis cycle.

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The difference between the observations and the background,

$$y_o - H(\mathbf{x}_b),$$

is called the **observational increment** or **innovation**.

The analysis \mathbf{x}_a is obtained by **adding the innovations to the background field** with **weights \mathbf{W}** that are determined based on the estimated statistical error covariances of the forecast and the observations:

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Earlier methods such as the SCM used weights which were determined **empirically**.

The weights were a function of the distance between the observation and the grid point, and the analysis was iterated several times.

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Lorenc (1986) showed that OI and the 3D-Var approach are equivalent if the cost function is defined as:

$$J = \frac{1}{2} \left\{ [\mathbf{y}_o - H(\mathbf{x})]^T \mathbf{R}^{-1} [\mathbf{y}_o - H(\mathbf{x})] + (\mathbf{x} - \mathbf{x}_b)^T \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}_b) \right\}$$

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The cost function J measures:

- The distance of a field \mathbf{x} to the observations (**first term**)
- The distance to the background \mathbf{x}_b (**second term**).

The distances are scaled by the observation error covariance R and by the background error covariance B respectively.

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- In OI, the weights \mathbf{W} are obtained for each grid point or grid volume, using suitable simplifications.
- In 3D-Var, the minimization of J is performed directly, allowing for additional flexibility and a simultaneous global use of the data.

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However, why should an observation in New Zealand be used to determine the pressure pattern in Ireland?

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The coefficients were determined by minimizing the mean square difference

$$\min_{a_{ij}} E = \min_{a_{ij}} \left[\sum_{k=1}^{K_z} p_k (z_k^o - z(x_k, y_k))^2 + \sum_{k=1}^{K_v} q_k \left\{ [u_k^o - u_g(x_k, y_k)]^2 + [v_k^o - v_g(x_k, y_k)]^2 \right\} \right]$$

Here p_k, q_k are empirical weighting coefficients and K is the total number of observations within the radius of influence.

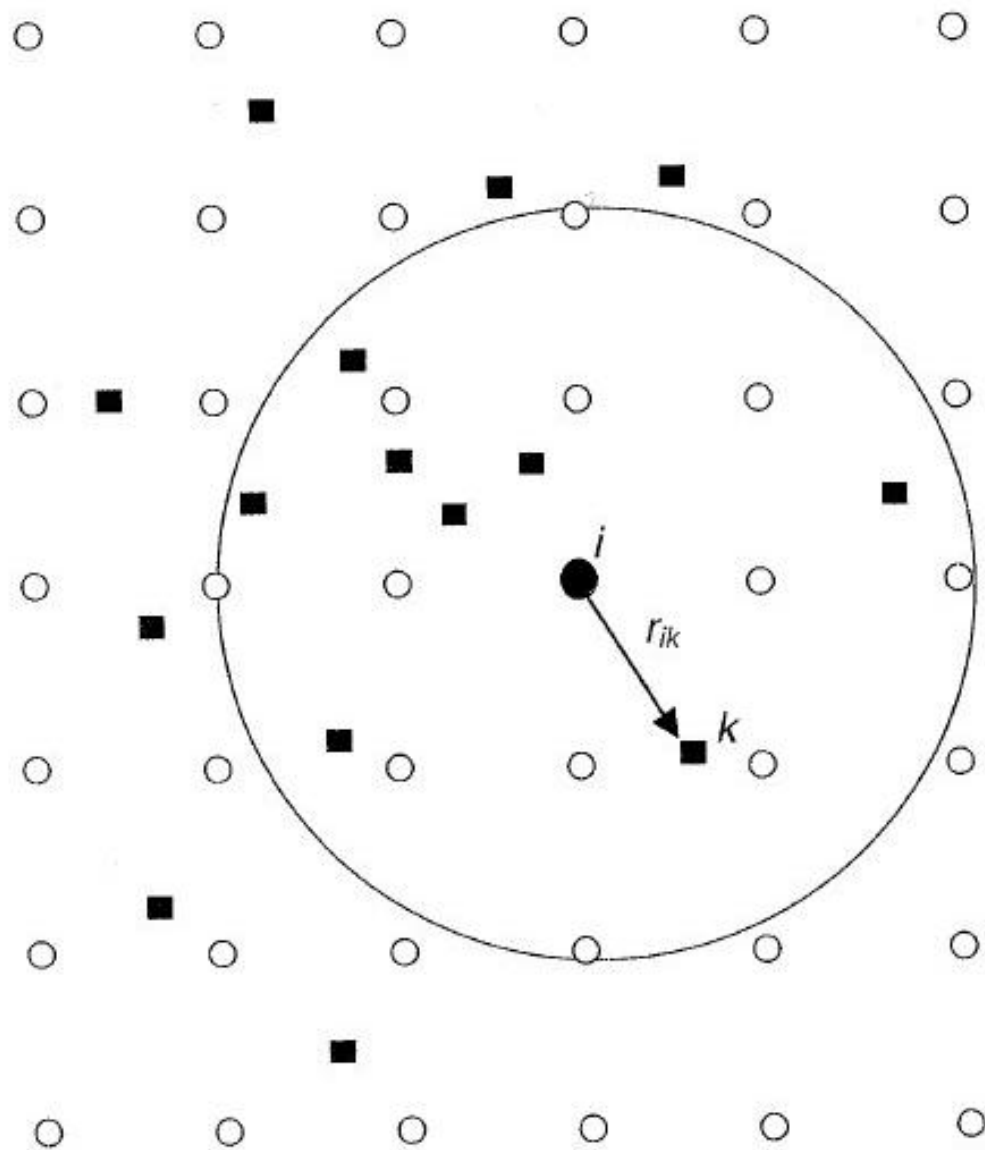


Figure 5.1.1: Schematic of grid points (circles), irregularly distributed observations (squares), and a radius of influence around a grid point i marked with a black circle. In 4DDA, the grid-point analysis is a combination of the forecast at the grid point (first guess) and the observational increments (observation minus first guess) computed at the observational points k . In certain analysis schemes, like SCM, only observations within the radius of influence, indicated by a circle, affect the analysis at the black grid point.

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Exercise: Consider the Gilchrist and Cressman scheme. What does the analysis look like if there is (i) a single pressure observation; (ii) two pressure observations close together; (iii) two pressure obs. far apart?

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There are many new types of data, such as **satellite** and **radar** observations, but:

- they don't measure the variables used in the models
- their distribution in space and time is very nonuniform.

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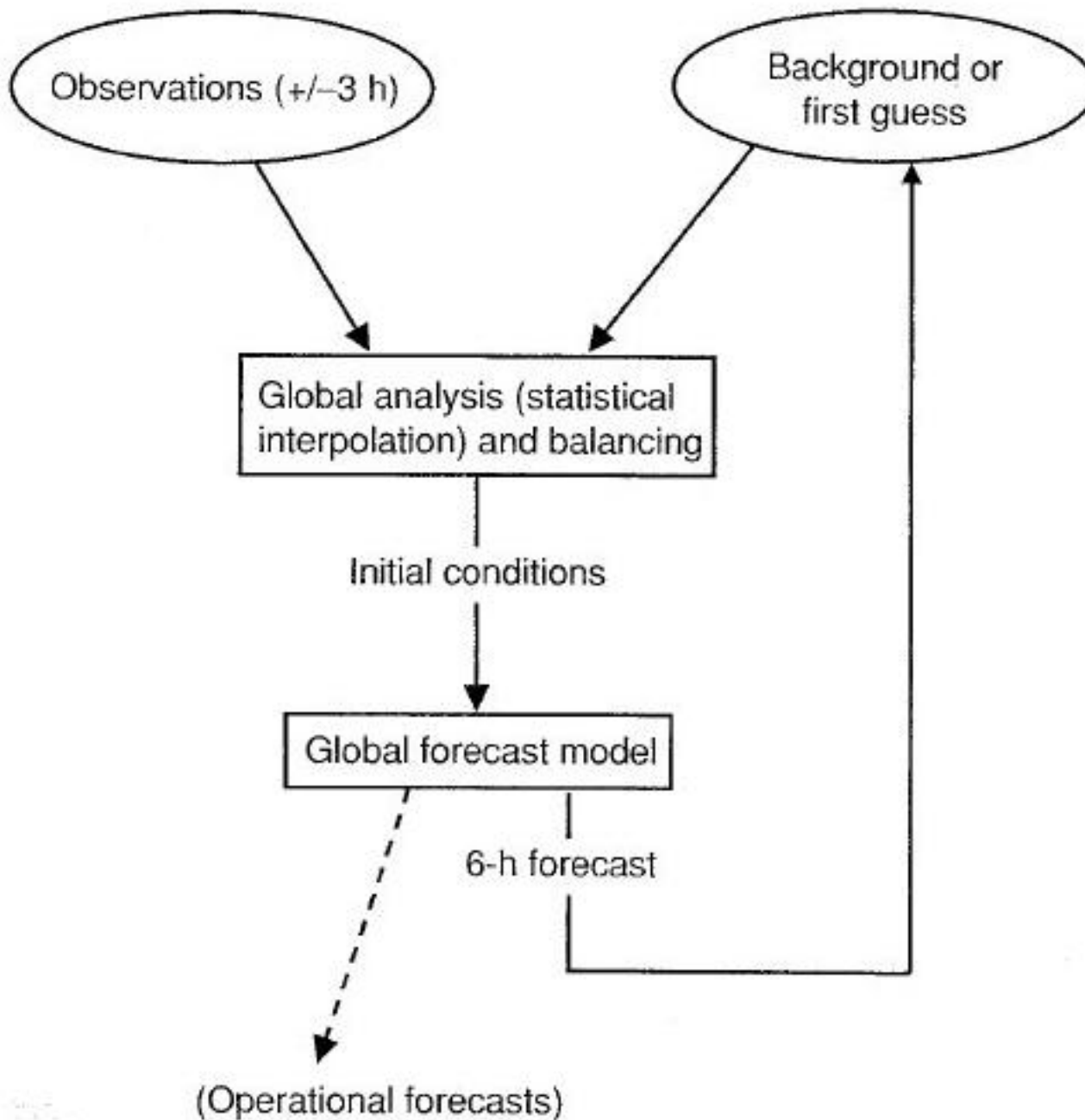
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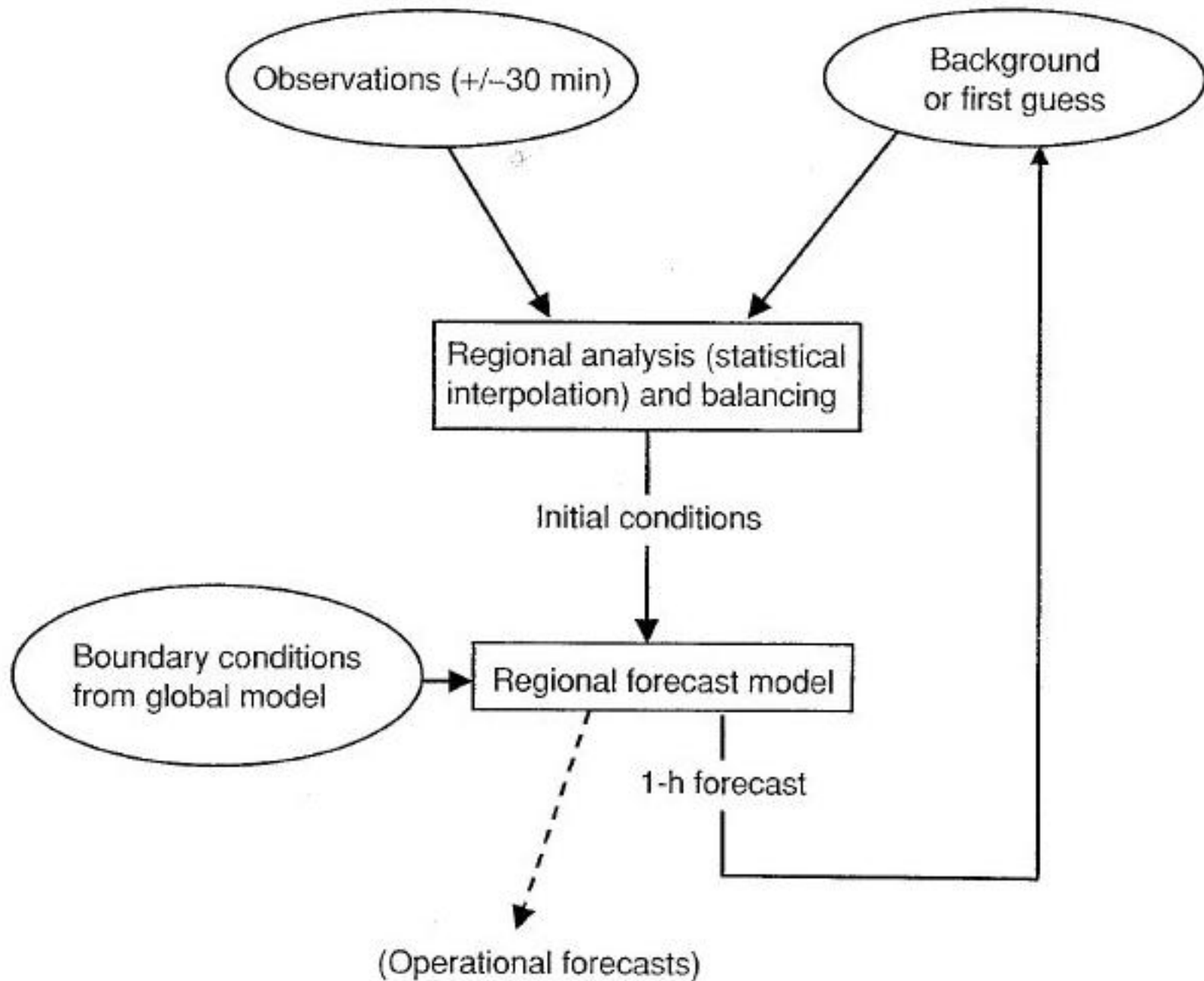
A short-range forecast is normally used as a first guess in operational systems in what is called an **analysis cycle**.

If a forecast is unavailable (e.g., if the cycle is broken), we may have to use **climatological fields** ...

... but they are normally a poor estimate of the initial state.



Global 6-h analysis cycle (00, 06, 12, and 18 UTC).



Regional analysis cycle, performed (perhaps) every hour.

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The model is able to **transport information** from data-rich to data-poor areas.

Exercise: Simple chart analysis.

