

M.Sc. in Meteorology

Synoptic Meteorology

[MAPH P312]

Prof Peter Lynch

Second Semester, 2004–2005

Seminar Room

Dept. of Maths. Physics, UCD, Belfield.

Part 9

Extratropical Weather Systems

These lectures follow closely the text of
Wallace & Hobbs (Chapter 9).

Introduction

We have studied the formation of raindrops by cloud microphysical processes. However, rainfall does not depend on microphysical processes alone.

Introduction

We have studied the formation of raindrops by cloud microphysical processes. However, rainfall does not depend on microphysical processes alone.

Without *organized large-scale upward motions*, the the hydrological cycle would stagnate:

- Tropospheric relative humidities would approach 100%
- Cloudiness would be widespread
- Precipitation would be minimal.

Introduction

We have studied the formation of raindrops by cloud microphysical processes. However, rainfall does not depend on microphysical processes alone.

Without *organized large-scale upward motions*, the the hydrological cycle would stagnate:

- Tropospheric relative humidities would approach 100%
- Cloudiness would be widespread
- Precipitation would be minimal.

Much of the ascent that drives the hydrological cycle in the earth's atmosphere occurs in association with *synoptic weather systems* with *well defined structures and life cycles*.

Introduction

We have studied the formation of raindrops by cloud microphysical processes. However, rainfall does not depend on microphysical processes alone.

Without *organized large-scale upward motions*, the hydrological cycle would stagnate:

- Tropospheric relative humidities would approach 100%
- Cloudiness would be widespread
- Precipitation would be minimal.

Much of the ascent that drives the hydrological cycle in the earth's atmosphere occurs in association with *synoptic weather systems* with *well defined structures and life cycles*.

We are concerned here with large scale extratropical weather systems — baroclinic waves and the associated extratropical cyclones — and with their embedded meso-scale fronts.

Extratropical Weather Systems

Day-to-day weather changes in midlatitudes are dominated by the passage of *baroclinic waves*, and their attendant

- cyclones
- fronts
- precipitation bands.

Extratropical Weather Systems

Day-to-day weather changes in midlatitudes are dominated by the passage of *baroclinic waves*, and their attendant

- cyclones
- fronts
- precipitation bands.

These large scale weather systems may assume a wide variety of forms, depending upon

- The background flow in which they are embedded
- The characteristics of the underlying surface
- The availability of moisture.

Extratropical Weather Systems

Day-to-day weather changes in midlatitudes are dominated by the passage of *baroclinic waves*, and their attendant

- cyclones
- fronts
- precipitation bands.

These large scale weather systems may assume a wide variety of forms, depending upon

- The background flow in which they are embedded
- The characteristics of the underlying surface
- The availability of moisture.

We will show how atmospheric data are analyzed to reveal the *structure* and *evolution* of weather systems.

Storm of 10th November, 1998

We will present a *case study* of a system that brought strong winds and heavy precipitation to parts of the central United States on November 10, 1998.

Storm of 10th November, 1998

We will present a *case study* of a system that brought strong winds and heavy precipitation to parts of the central United States on November 10, 1998.

This particular storm system was unusually strong, but it typifies many of the features of the life cycle of baroclinic waves and the associated extratropical cyclones.

Storm of 10th November, 1998

We will present a *case study* of a system that brought strong winds and heavy precipitation to parts of the central United States on November 10, 1998.

This particular storm system was unusually strong, but it typifies many of the features of the life cycle of baroclinic waves and the associated extratropical cyclones.

Although the storm occurred over the continental USA, it has all the important features of storms which occur in our neighbourhood: it is quite representative of storms which we see in Ireland.

Overview

We will examine and document the large scale structure of the November 10, 1998 storm, with emphasis on the

- 500-hPa height
- Sea-level pressure
- 1000–500-hPa thickness
- Vertical velocity fields.

Overview

We will examine and document the large scale structure of the November 10, 1998 storm, with emphasis on the

- 500-hPa height
- Sea-level pressure
- 1000–500-hPa thickness
- Vertical velocity fields.

The development of the storm is shown to be linked to the *intensification of a baroclinic wave* as it evolves through its life cycle.

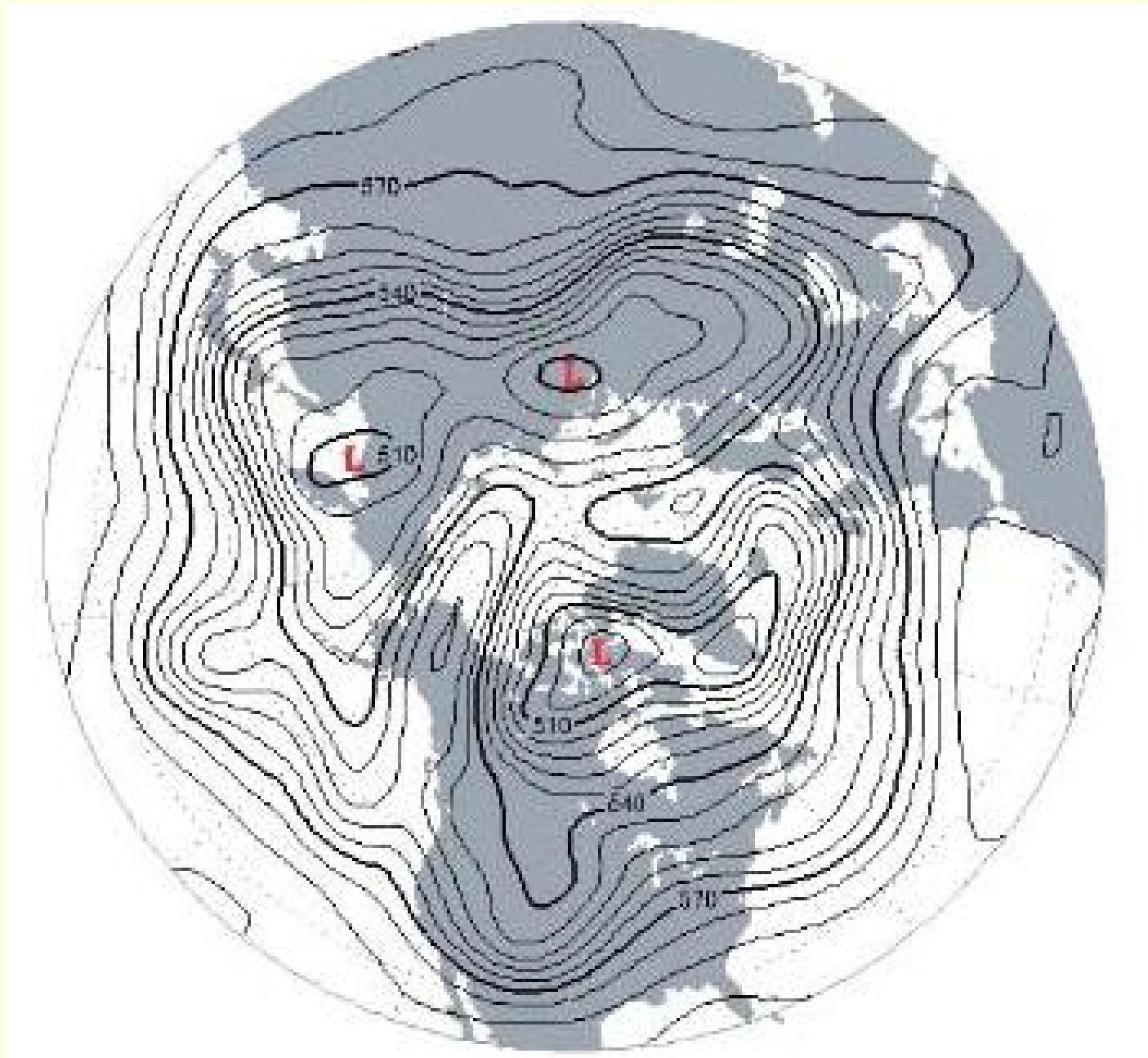
Overview

We will examine and document the large scale structure of the November 10, 1998 storm, with emphasis on the

- 500-hPa height
- Sea-level pressure
- 1000–500-hPa thickness
- Vertical velocity fields.

The development of the storm is shown to be linked to the *intensification of a baroclinic wave* as it evolves through its life cycle.

The hemispheric 500-hPa chart for Midnight Universal Time (0000 UTC), November 10, 1998 is shown below.



500-hPa height chart for 00 Nov. 10, 1998. Contour interval 60 m.

The basic structure is a **circumpolar cyclonic vortex**. However, the vortex is *far from axisymmetric*.

The basic structure is a **circumpolar cyclonic vortex**. However, the vortex is *far from axisymmetric*.

A pair of ridges, one extending up over Alaska and the other through Scandinavia and Spitzbergen, have temporarily split the westerly **polar vortex** into more regional vortices, one centered over Russia and the other centered over Northern Canada.

The basic structure is a **circumpolar cyclonic vortex**. However, the vortex is *far from axisymmetric*.

A pair of ridges, one extending up over Alaska and the other through Scandinavia and Spitzbergen, have temporarily split the westerly **polar vortex** into more regional vortices, one centered over Russia and the other centered over Northern Canada.

These **blocking ridges** are notable for their persistence, which often exceeds a week.

The basic structure is a **circumpolar cyclonic vortex**. However, the vortex is *far from axisymmetric*.

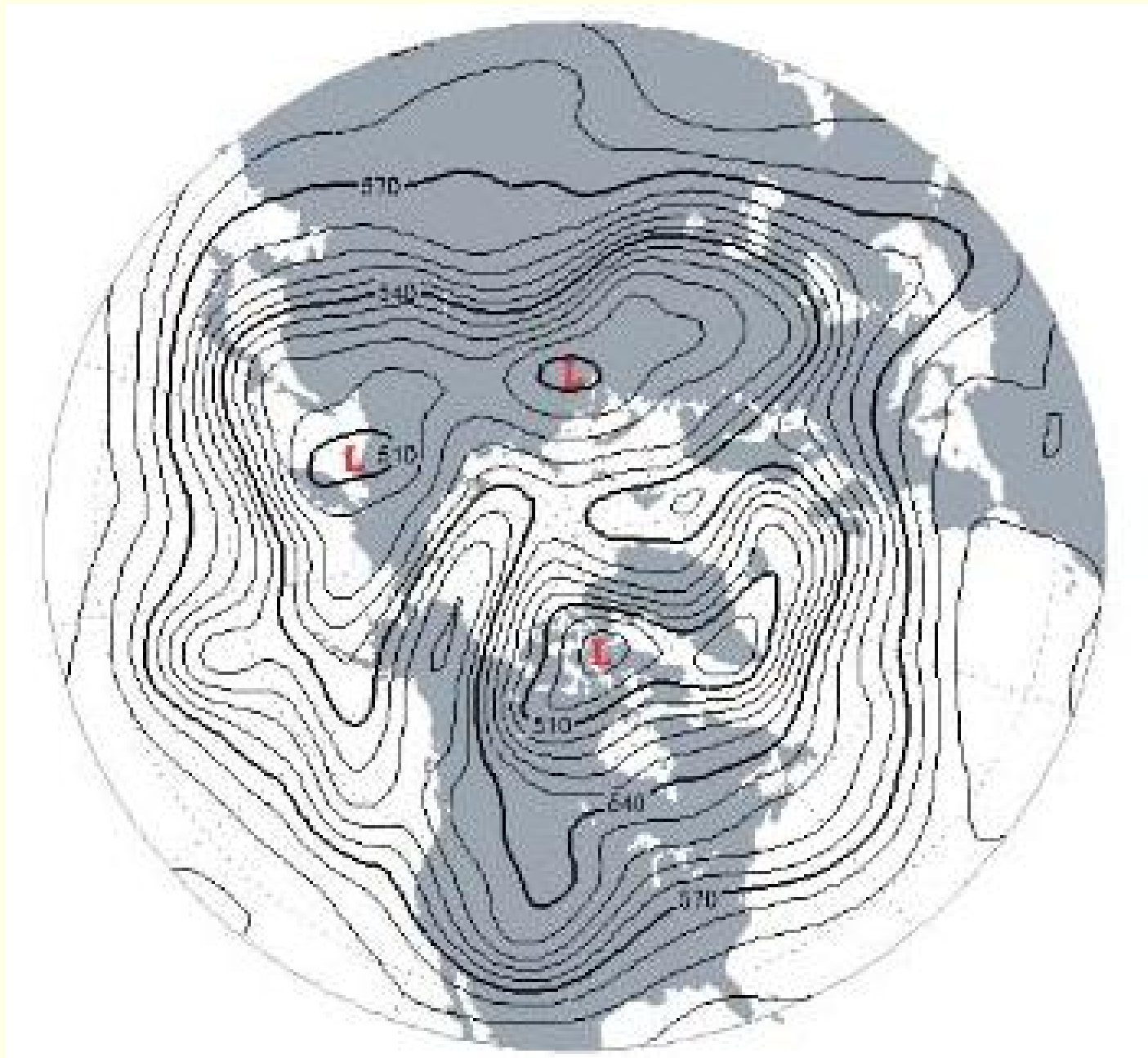
A pair of ridges, one extending up over Alaska and the other through Scandinavia and Spitzbergen, have temporarily split the westerly **polar vortex** into more regional vortices, one centered over Russia and the other centered over Northern Canada.

These **blocking ridges** are notable for their persistence, which often exceeds a week.

Pronounced troughs are evident over

- The Black Sea
- Japan
- The central Pacific
- the United States Great Plains

with several weaker troughs at other locations.



500-hPa height chart for 00 Nov. 10, 1998. Contour interval 60 m.

There are about **eight troughs** (counting the weaker ones) around the parallel of latitude at 45° North.

Thus, the typical distance between successive troughs is

$$\frac{360^\circ}{8} = 45^\circ \quad \text{or about } 45 \text{ degrees of longitude.}$$

There are about **eight troughs** (counting the weaker ones) around the parallel of latitude at 45° North.

Thus, the typical distance between successive troughs is

$$\frac{360^\circ}{8} = 45^\circ \quad \text{or about } 45 \text{ degrees of longitude.}$$

The length of the parallel at 45°N is

$$2\pi a \cos 45^\circ = 4 \times 10^7 \times \cos 45^\circ \approx 28,000 \text{ km}$$

So, the wavelength at 45°N is one-eighth of this:

$$L(45^\circ\text{N}) = \frac{28,000}{8} \approx 3500 \text{ km}$$

The scale of 3500 km corresponds approximately to the theoretically predicted ***wavelength of baroclinic waves***.

There are about **eight troughs** (counting the weaker ones) around the parallel of latitude at 45° North.

Thus, the typical distance between successive troughs is

$$\frac{360^\circ}{8} = 45^\circ \quad \text{or about 45 degrees of longitude.}$$

The length of the parallel at 45°N is

$$2\pi a \cos 45^\circ = 4 \times 10^7 \times \cos 45^\circ \approx 28,000 \text{ km}$$

So, the wavelength at 45°N is one-eighth of this:

$$L(45^\circ\text{N}) = \frac{28,000}{8} \approx 3500 \text{ km}$$

The scale of 3500 km corresponds approximately to the theoretically predicted ***wavelength of baroclinic waves***.

Baroclinic waves tend to propagate eastward with a phase speed on the order of 10 m s^{-1} , which corresponds to the wintertime climatological-mean zonal wind speed around the 700-hPa level.

The *strength of the westerlies* generally tends to increase with height within the extratropical troposphere.

The *strength of the westerlies* generally tends to increase with height within the extratropical troposphere.

Hence, above this so-called *steering level*, air parcels tend to pass through the waves from west to east.

Below the *steering level*, the waves propagate more rapidly than air parcels are typically moving.

The *strength of the westerlies* generally tends to increase with height within the extratropical troposphere.

Hence, above this so-called *steering level*, air parcels tend to pass through the waves from west to east.

Below the *steering level*, the waves propagate more rapidly than air parcels are typically moving.

Taking the wavelength as $L = 3500 \text{ km}$ and the phase speed as $c = 10 \text{ m s}^{-1}$, the period is

$$\tau = \frac{L}{c} = \frac{3.5 \times 10^6}{10} = 350,000 \text{ s} \approx 4 \text{ days}$$

The *strength of the westerlies* generally tends to increase with height within the extratropical troposphere.

Hence, above this so-called *steering level*, air parcels tend to pass through the waves from west to east.

Below the *steering level*, the waves propagate more rapidly than air parcels are typically moving.

Taking the wavelength as $L = 3500 \text{ km}$ and the phase speed as $c = 10 \text{ m s}^{-1}$, the period is

$$\tau = \frac{L}{c} = \frac{3.5 \times 10^6}{10} = 350,000 \text{ s} \approx 4 \text{ days}$$

Depending on the strength of the steering flow, successive ridges (or troughs) typically pass a fixed point on Earth at intervals around 4 days apart.

The *strength of the westerlies* generally tends to increase with height within the extratropical troposphere.

Hence, above this so-called *steering level*, air parcels tend to pass through the waves from west to east.

Below the *steering level*, the waves propagate more rapidly than air parcels are typically moving.

Taking the wavelength as $L = 3500$ km and the phase speed as $c = 10 \text{ m s}^{-1}$, the period is

$$\tau = \frac{L}{c} = \frac{3.5 \times 10^6}{10} = 350,000 \text{ s} \approx 4 \text{ days}$$

Depending on the strength of the steering flow, successive ridges (or troughs) typically pass a fixed point on Earth at intervals around 4 days apart.

The interval may be *as short as a day or two* if the steering flow is very strong, and it may be *a week or longer* if the westerlies aloft are interrupted by **blocking**.

The direction of propagation is usually **eastward**, but it may may be **northeastward** or **southeastward** if the steering flow meanders.

The direction of propagation is usually **eastward**, but it may may be **northeastward** or **southeastward** if the steering flow meanders.

Baroclinic waves are observed most regularly and tend to be strongest *over the oceans*, but they also develop over land, as in this case study.

The direction of propagation is usually **eastward**, but it may may be **northeastward** or **southeastward** if the steering flow meanders.

Baroclinic waves are observed most regularly and tend to be strongest *over the oceans*, but they also develop over land, as in this case study.

The storms which affect Ireland frequently form off the eastern seaboard of USA and Canada.

The direction of propagation is usually **eastward**, but it may may be **northeastward** or **southeastward** if the steering flow meanders.

Baroclinic waves are observed most regularly and tend to be strongest *over the oceans*, but they also develop over land, as in this case study.

The storms which affect Ireland frequently form off the eastern seaboard of USA and Canada.

They *tend to be* most vigorous during the winter months of the year, when the polar regions are cold and the meridional temperature gradient across mid-latitudes is strongest.

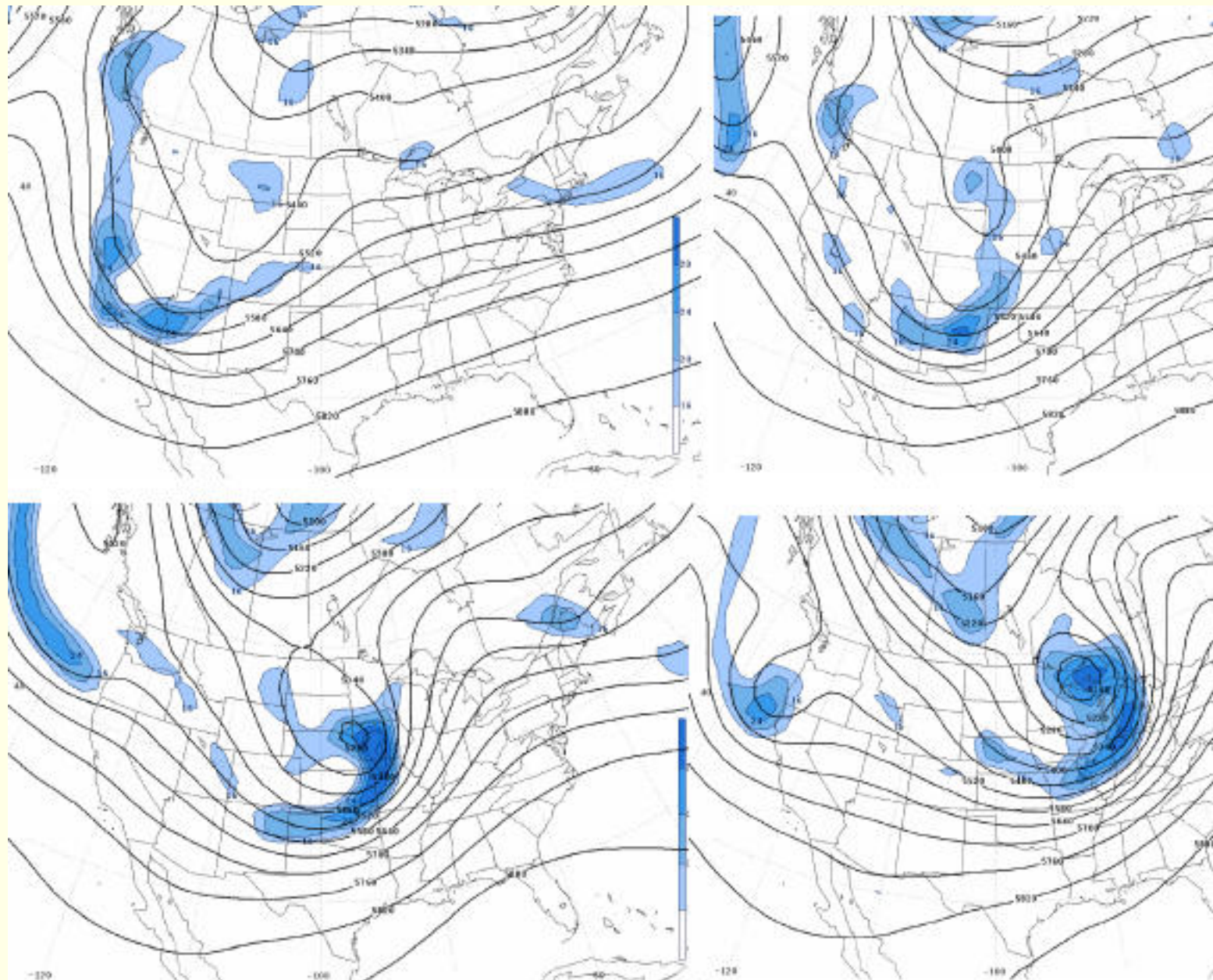
The direction of propagation is usually **eastward**, but it may may be **northeastward** or **southeastward** if the steering flow meanders.

Baroclinic waves are observed most regularly and tend to be strongest *over the oceans*, but they also develop over land, as in this case study.

The storms which affect Ireland frequently form off the eastern seaboard of USA and Canada.

They *tend to be* most vigorous during the winter months of the year, when the polar regions are cold and the meridional temperature gradient across mid-latitudes is strongest.

A more detailed view of the 500-hPa height pattern for 00 UTC, 10 November over the North American sector is shown below, together with charts for 12 hours earlier and 12 and 24 hours later.

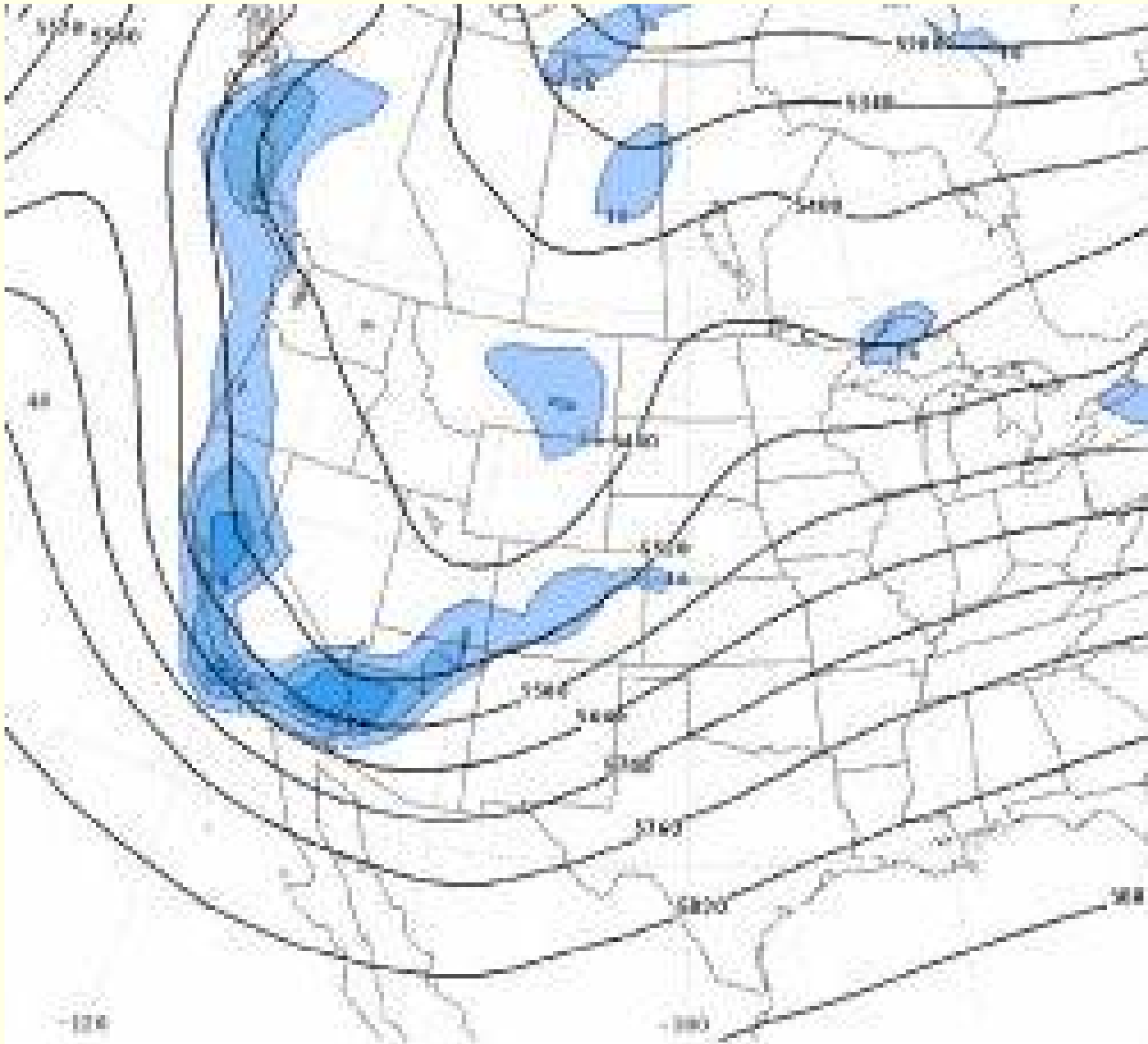


500-hPa height and absolute vorticity (coloured shading) at 12 hour intervals, starting at 12 UTC Nov. 9, 1998.

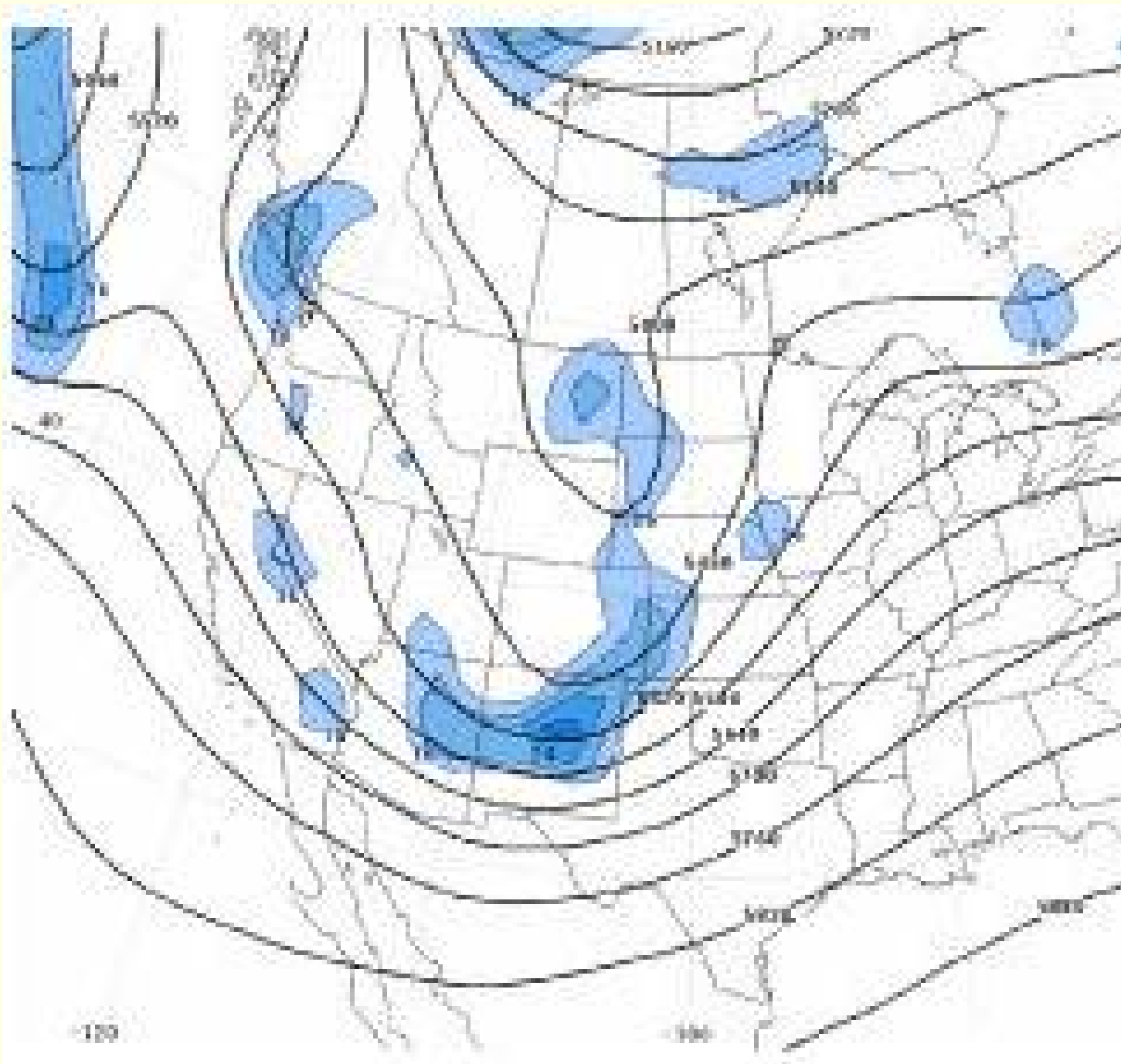
Clearly evident in this 4-chart sequence are the *eastward propagation* and *intensification* of the trough that passes over the Great Plains and the ridge immediately downstream of it.

Clearly evident in this 4-chart sequence are the *eastward propagation* and *intensification* of the trough that passes over the Great Plains and the ridge immediately downstream of it.

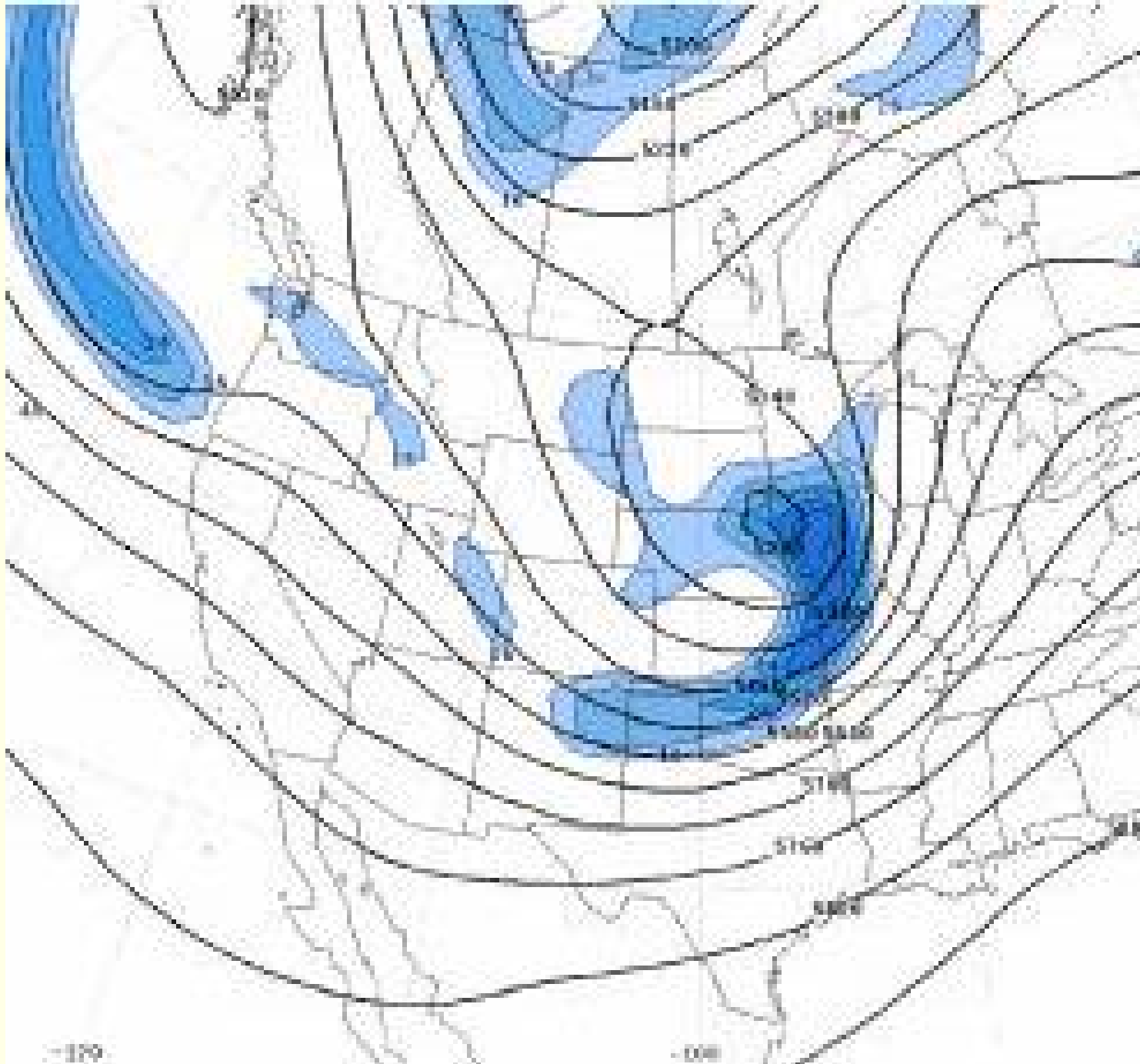
Let us look in more detail at the sequence. The contours are 500-hPa height and the coloured areas are regions with high values of absolute vorticity ($f + \zeta$).



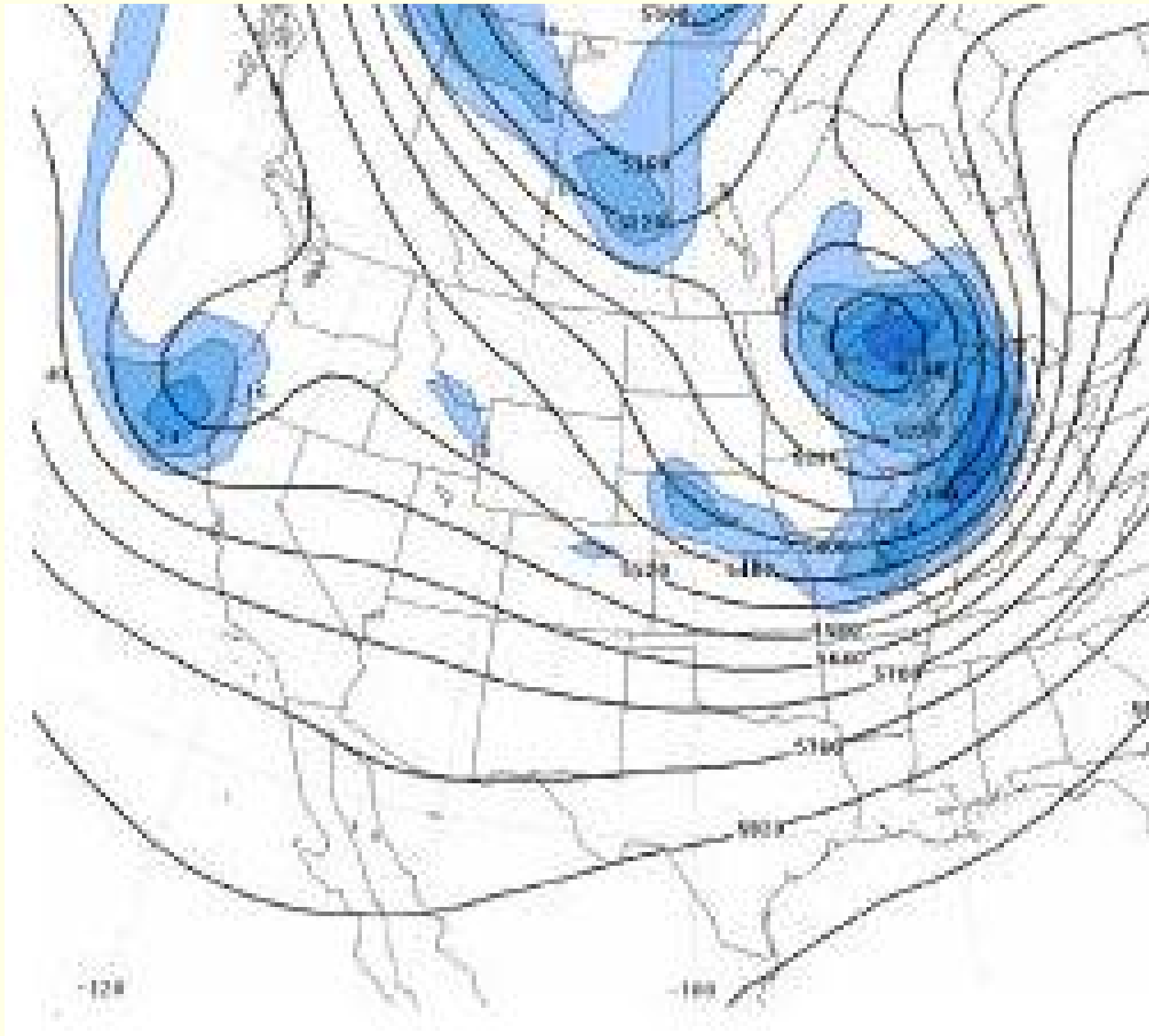
12 UTC, 9 Nov., 1998.



00 UTC, 10 Nov., 1998.



12 UTC, 10 Nov., 1998.



In the third and fourth charts in the sequence the base of this trough splits off to form a cutoff low, i.e., an isolated minimum in the geopotential height field, implying the existence of a **closed cyclonic circulation**.

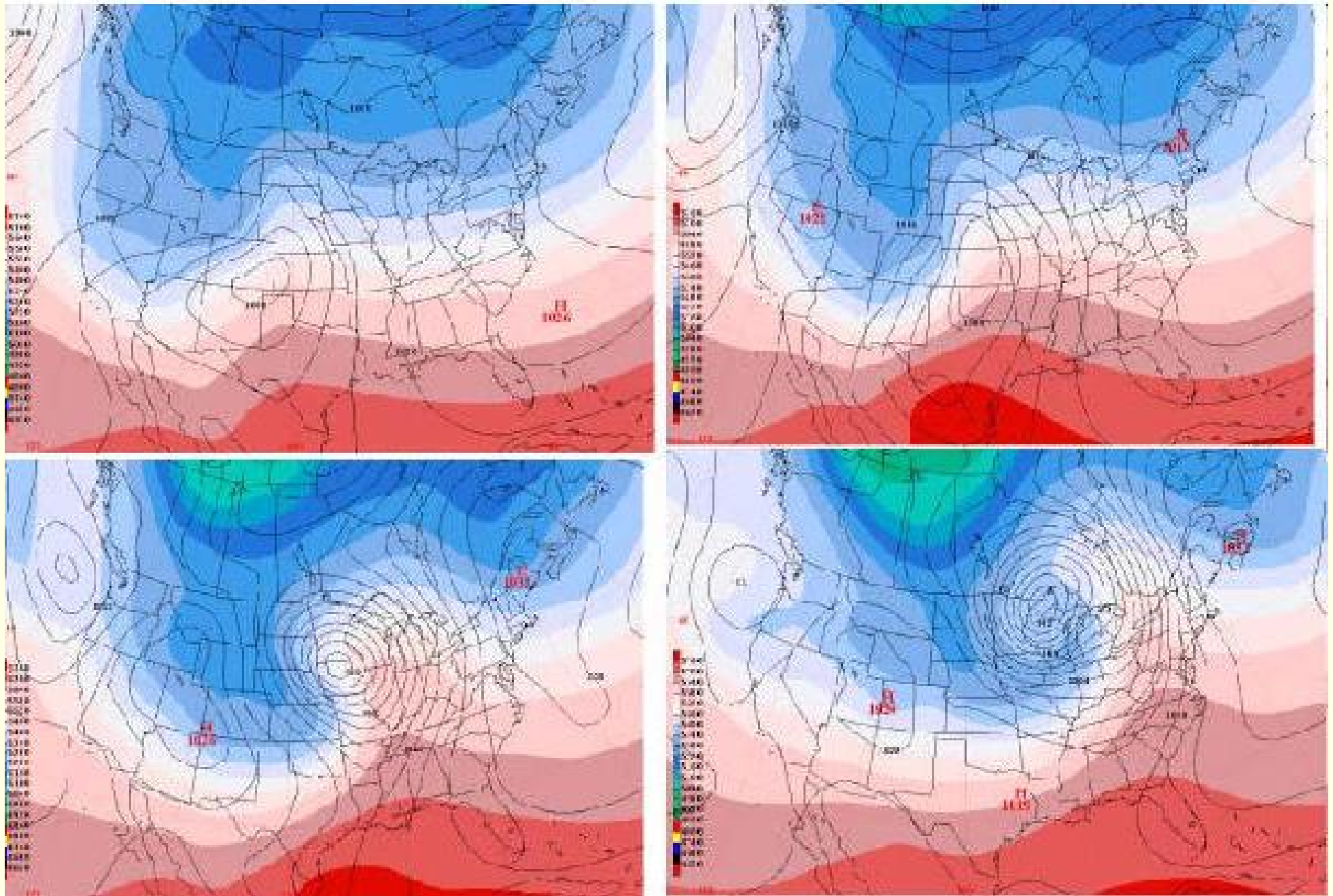
In the third and fourth charts in the sequence the base of this trough splits off to form a cutoff low, i.e., an isolated minimum in the geopotential height field, implying the existence of a **closed cyclonic circulation**.

The winds encircling this developing feature strengthen, as evidenced by the tightening of the spacing between adjacent 500-hPa height contours.

In the third and fourth charts in the sequence the base of this trough splits off to form a cutoff low, i.e., an isolated minimum in the geopotential height field, implying the existence of a **closed cyclonic circulation**.

The winds encircling this developing feature strengthen, as evidenced by the tightening of the spacing between adjacent 500-hPa height contours.

The intensification of the trough at the 500-hPa level is accompanied by the deepening of the corresponding low pressure center in sea-level pressure field (shown below).



Sea-level pressure (contours) and 1000–500 hPa thickness (shading).

This surface low marks the center of a closed cyclonic circulation referred to as an extratropical cyclone.

This surface low marks the center of a closed cyclonic circulation referred to as an extratropical cyclone.

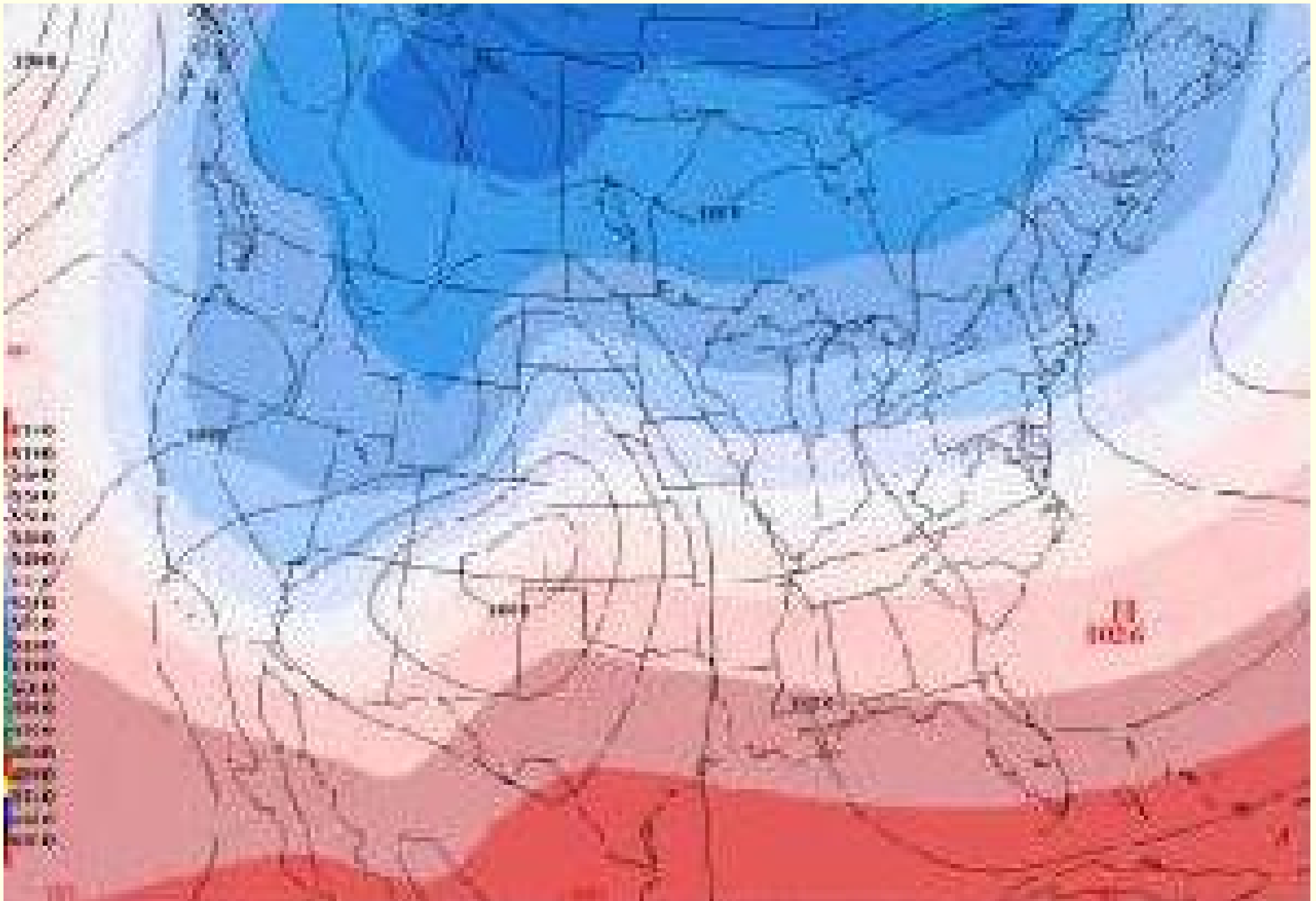
Also evident is the amplification of the east-west gradients in the 1000–500 hPa thickness field, indicated by the colored shading.

This surface low marks the center of a closed cyclonic circulation referred to as an **extratropical cyclone**.

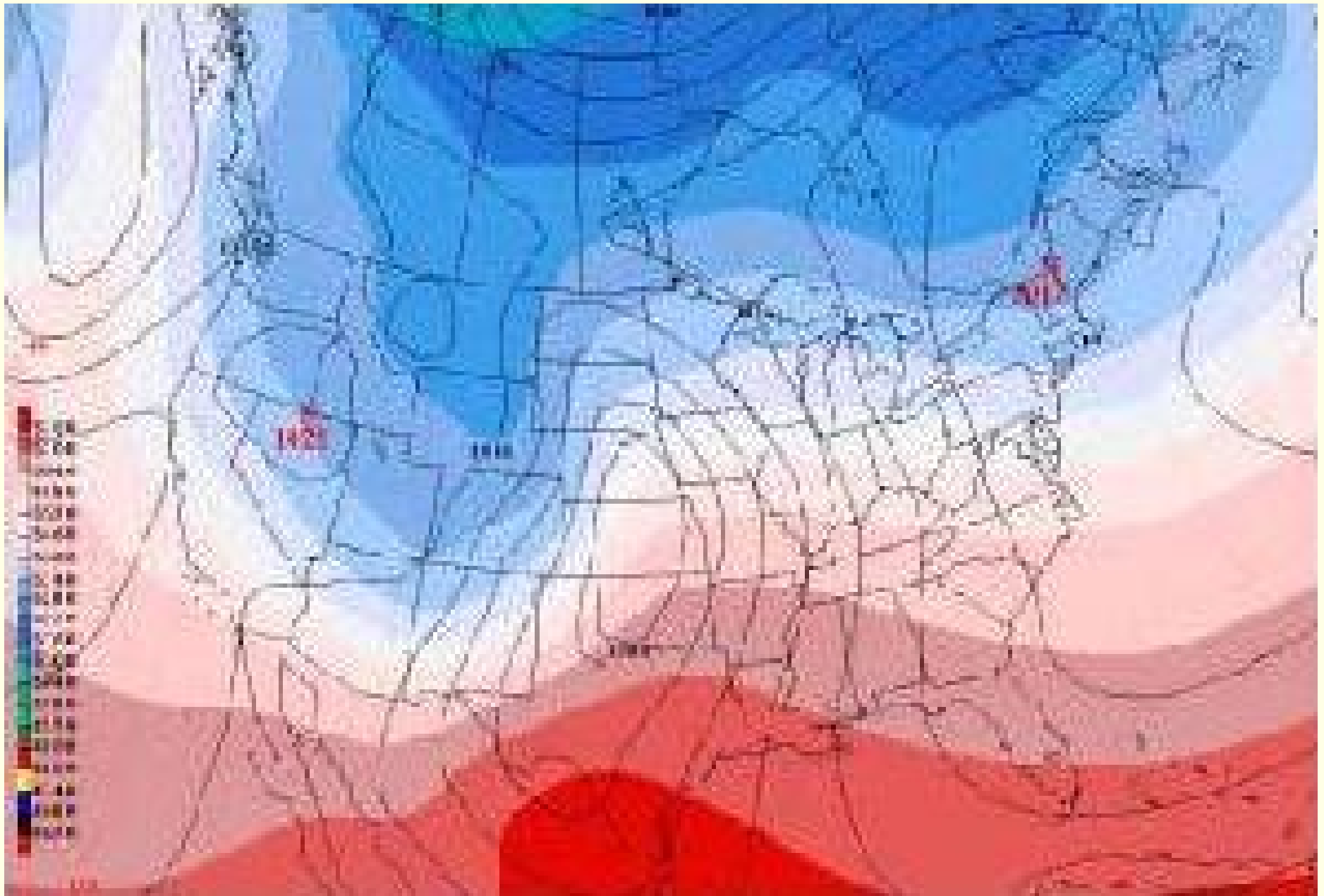
Also evident is the amplification of the east-west gradients in the **1000–500 hPa thickness field**, indicated by the colored shading.

In the first two charts of the sequence the developing surface low is located well to the east of the corresponding trough in the 500-hPa height field.

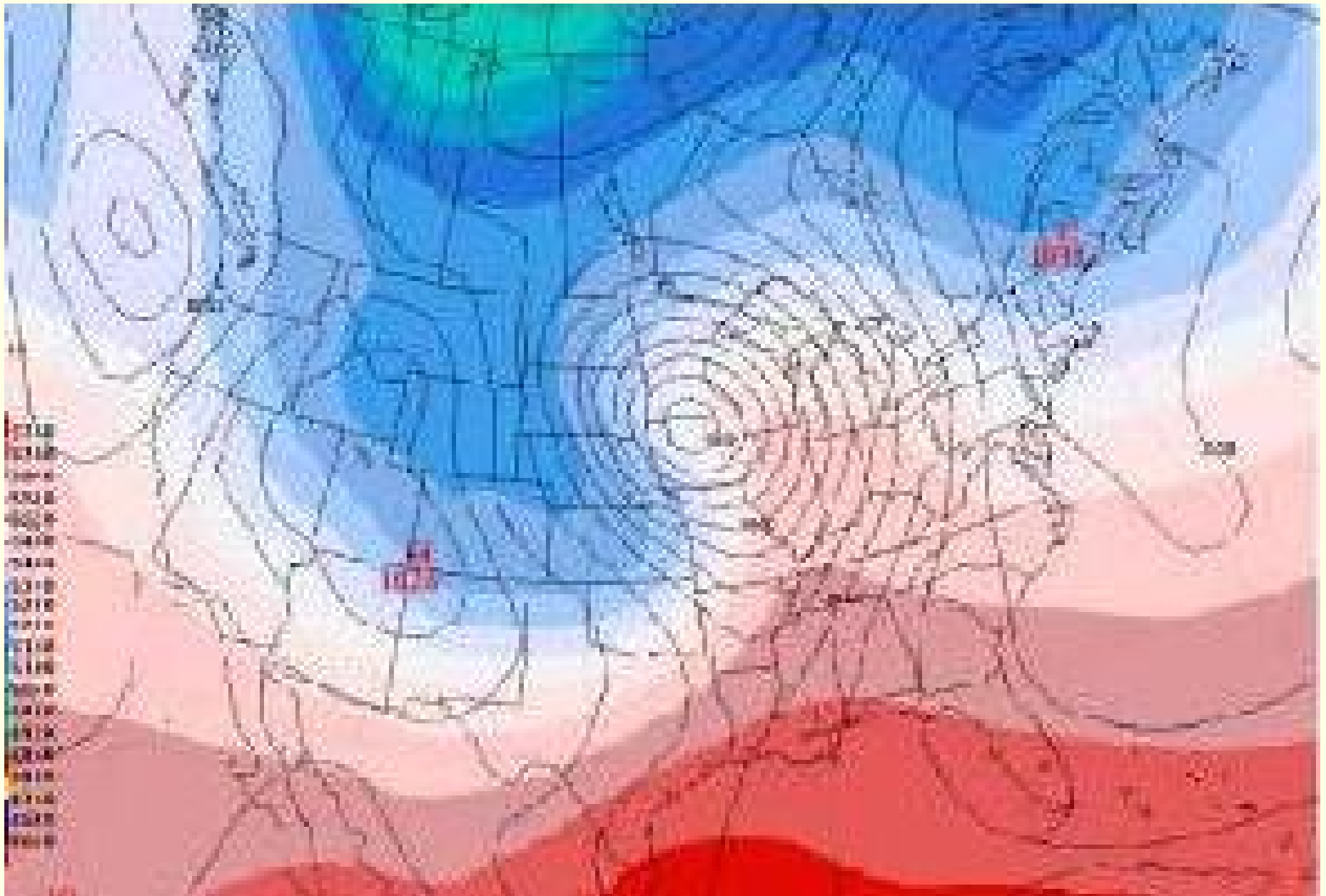
As these features amplify, they come into vertical alignment in the later charts of the sequence.



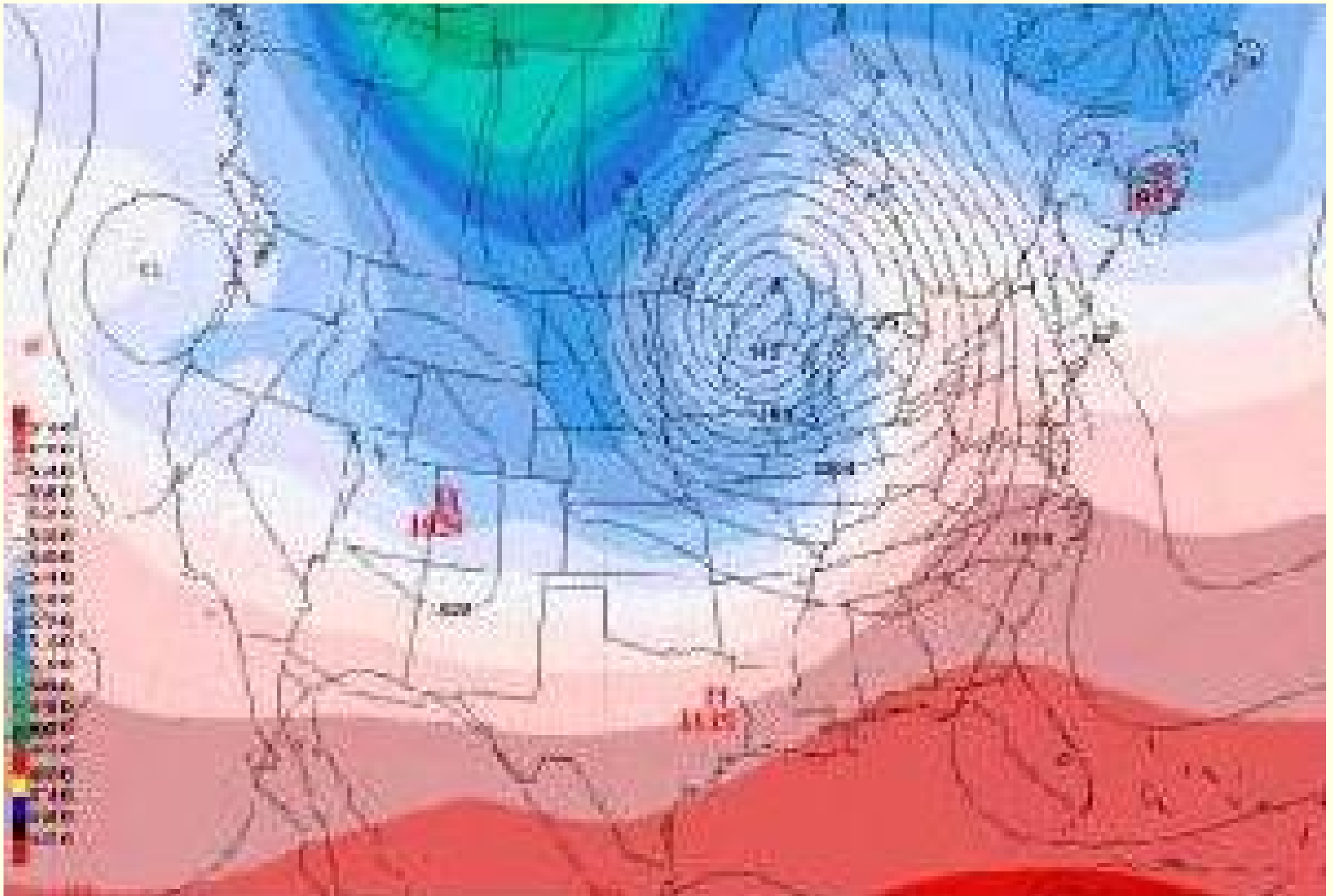
12 UTC, 9 Nov., 1998.



00 UTC, 10 Nov., 1998.



12 UTC, 10 Nov., 1998.



00 UTC, 11 Nov., 1998.