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$$\left[\nabla^2 + \frac{\partial}{\partial p} \left(\frac{f_0^2}{\sigma} \frac{\partial}{\partial p} \right) \right] \Phi_t = - f_0 \mathbf{V}_g \cdot \nabla \left(\frac{1}{f_0} \nabla^2 \Phi + f \right) + \frac{\partial}{\partial p} \left[\frac{f_0^2}{\sigma} \mathbf{V}_g \cdot \nabla \left(-\frac{\partial \Phi}{\partial p} \right) \right]$$

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The second term on the right (Term (C)) may be expanded:

$$-\mathbf{V}_g \cdot \nabla \frac{\partial}{\partial p} \left(\frac{f_0^2}{\sigma} \frac{\partial \Phi}{\partial p} \right) - \frac{f_0^2}{\sigma} \frac{\partial \mathbf{V}_g}{\partial p} \cdot \nabla \frac{\partial \Phi}{\partial p}$$

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But the thermal wind relationship is

$$f_0 \frac{\partial \mathbf{V}_g}{\partial p} = \mathbf{k} \times \nabla \frac{\partial \Phi}{\partial p}$$

This is just the p -derivative of $f_0 \mathbf{V}_g = \mathbf{k} \times \nabla \Phi$.

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The remaining term can be combined with term (B) in the tendency equation to give

$$\text{RHS} = -f_0 \mathbf{V}_g \cdot \nabla \left[\frac{1}{f_0} \nabla^2 \Phi + f + \frac{\partial}{\partial p} \left(\frac{f_0}{\sigma} \frac{\partial \Phi}{\partial p} \right) \right] = -f_0 \mathbf{V}_g \cdot \nabla q$$

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The quantity in square brackets is called the *quasi-geostrophic potential vorticity*

$$q \equiv \left[\frac{1}{f_0} \nabla^2 \Phi + f + \frac{\partial}{\partial p} \left(\frac{f_0}{\sigma} \frac{\partial \Phi}{\partial p} \right) \right]$$

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The left side of the tendency equation may be written

$$\text{LHS} = f_0 \frac{\partial}{\partial t} \left[\frac{1}{f_0} \nabla^2 \Phi + \frac{\partial}{\partial p} \left(\frac{f_0}{\sigma} \frac{\partial \Phi}{\partial p} \right) \right] = f_0 \frac{\partial q}{\partial t}$$

since f does not vary with time.

The tendency equation may now be written in a conservative form called the *quasi-geostrophic potential vorticity equation* or QGPV equation:

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Note that q is completely determined once the three-dimensional distribution of geopotential Φ is given.

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This was the equation

$$\frac{d}{dt}(\zeta + f) = 0$$

for the conservation of *absolute vorticity*.

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An idealized geopotential field is given at time $t = 0$ by

$$\Phi = \Phi_0 - f_0 \bar{u} y + A \sin(kx - mp)$$

where Φ_0 , \bar{u} and A are functions of p .

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- (a) Compute the geostrophic wind components as functions of x and p .
- (b) Ignoring the β -effect, compute the geostrophic vorticity and divergence.
- (c) Compute the variations in the temperature field due to the wave.

(d) Compute the vorticity advection and temperature advection.

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- (e) Using the geopotential tendency equation, describe how the pressure at a point upstream from a trough and downstream from a ridge is expected to change.

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- (e) Using the geopotential tendency equation, describe how the pressure at a point upstream from a trough and downstream from a ridge is expected to change.
- (f) Using the omega equation, describe the pattern of vertical velocity associated with the wave disturbance.

The ENIAC Integrations

(ENIAC: Electronic Numerical Integrator and Computer)

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1. Entirely new methods of weather prediction by calculation will have been made possible;
2. A new rational basis will have been secured for the planning of physical measurements and field observations;
3. The first step towards influencing the weather by rational human intervention will have been made.

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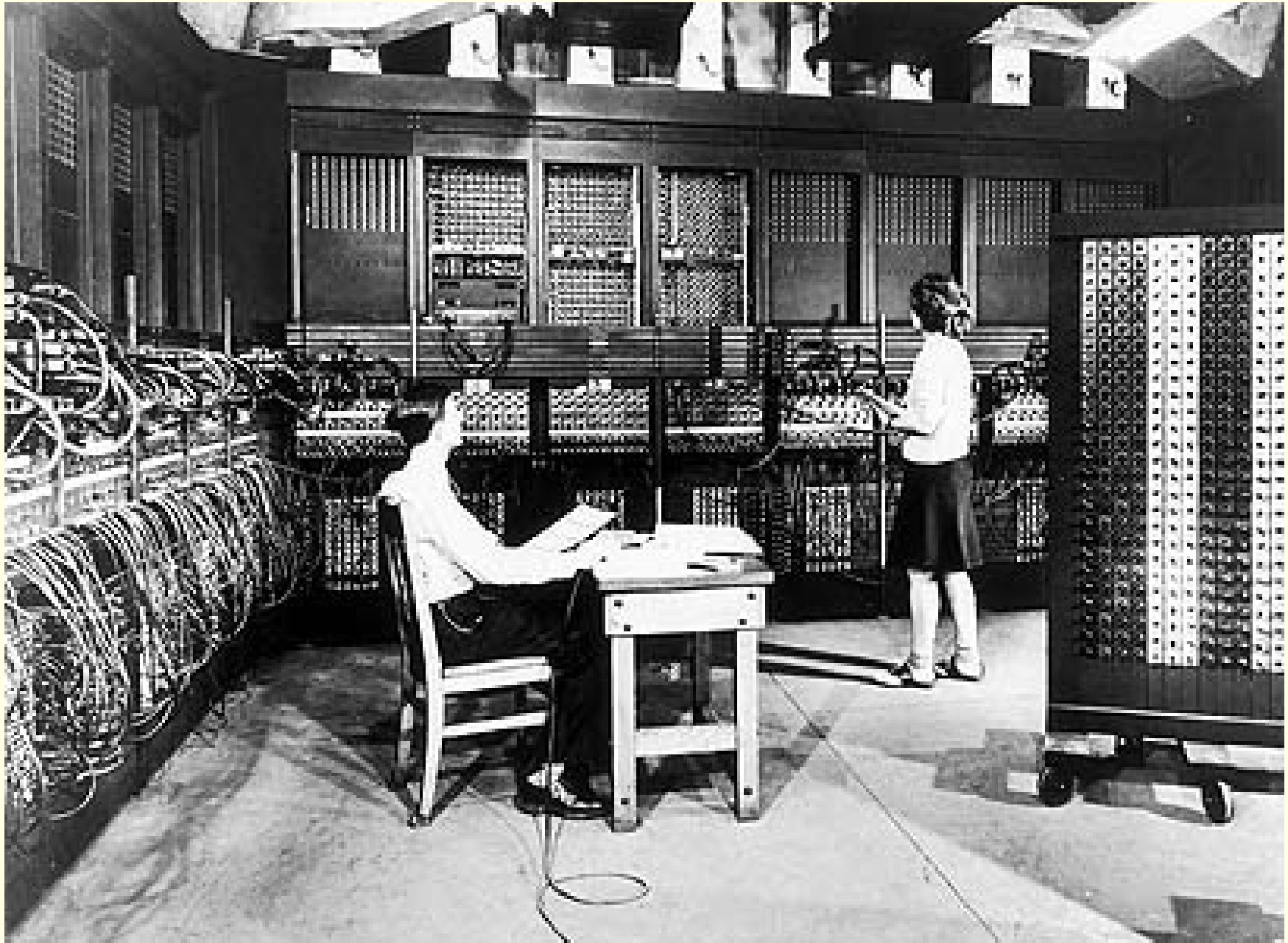
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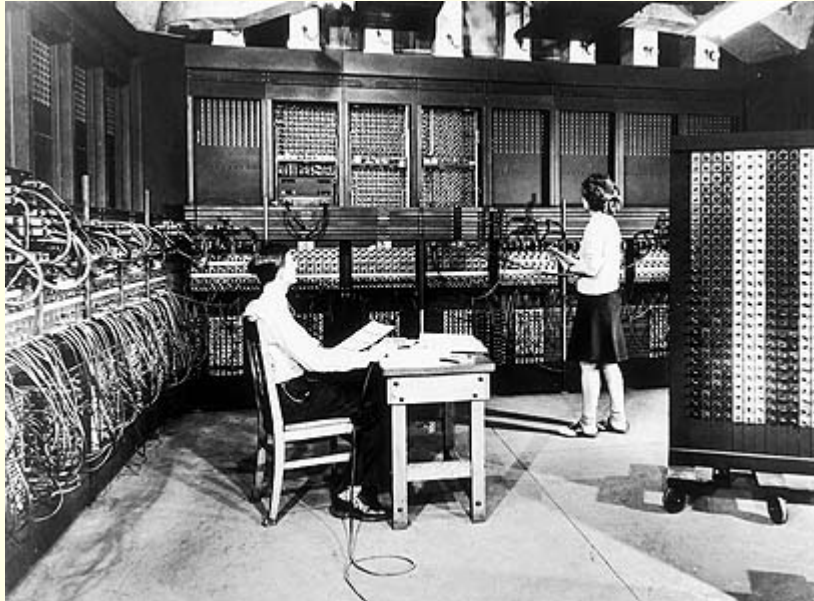
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The ENIAC



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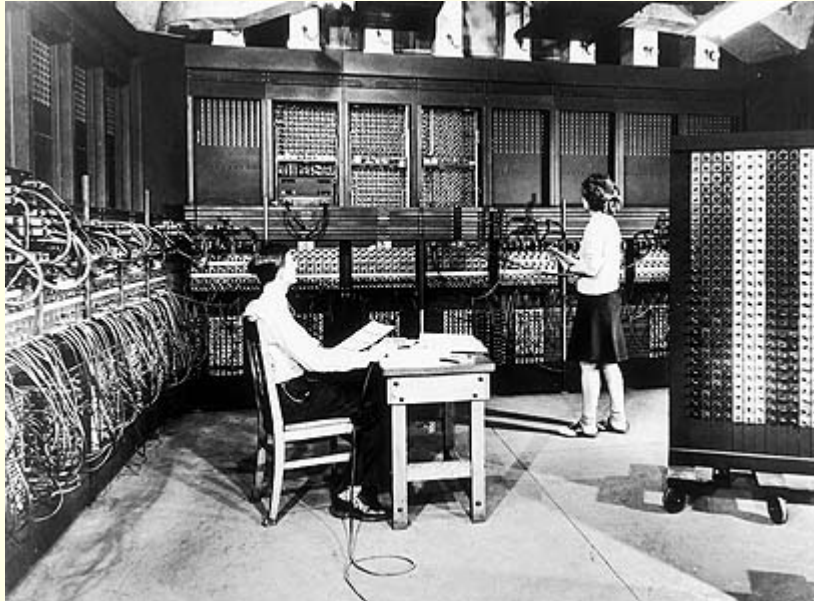


The **ENIAC** (Electronic Numerical Integrator and Computer) was the *first multi-purpose programmable electronic digital computer*.

It had:

- 18,000 vacuum tubes
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Power Consumption: 140 kWatts

The ENIAC: Technical Details.

ENIAC was a **decimal machine**. No high-level language.
Assembly language. Fixed-point arithmetic: $-1 < x < +1$.
10 registers, that is,

Ten words of high-speed memory.

Function Tables:

624 6-digit words of “ROM”, set on
ten-pole rotary switches.

“Peripheral Memory”:

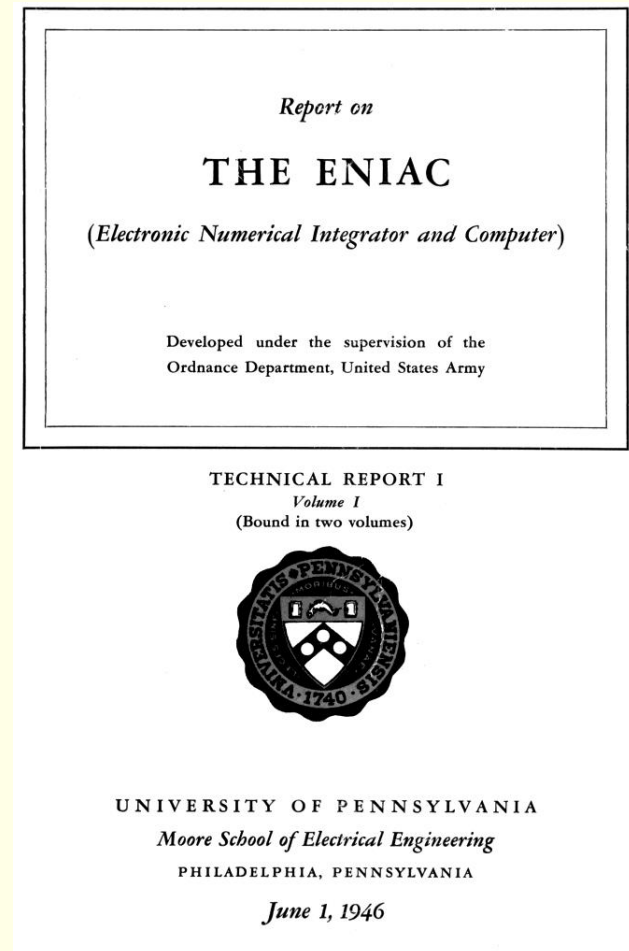
Punch-cards.

Speed: FP multiply: 2ms
(say, **500 Flops**).

Access to Function Tables: **1ms.**

Access to Punch-card equipment:

You can imagine!



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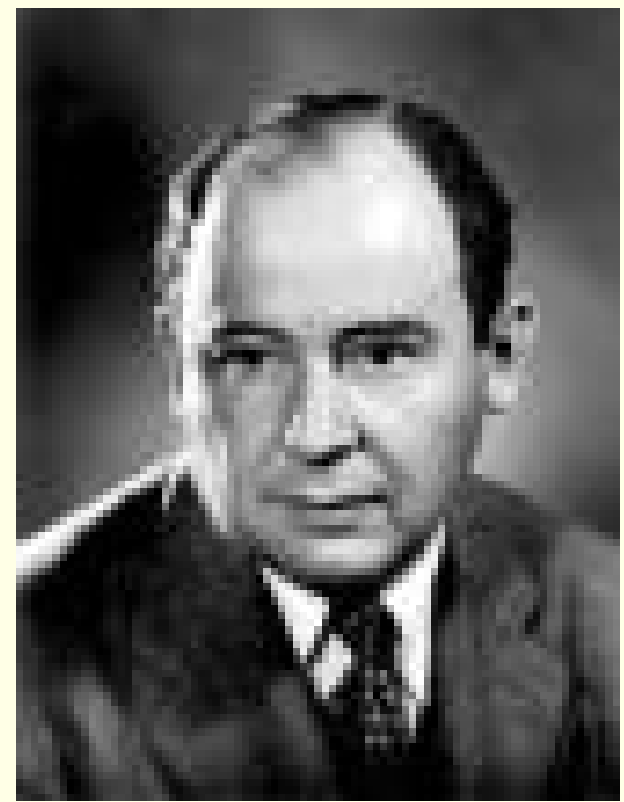
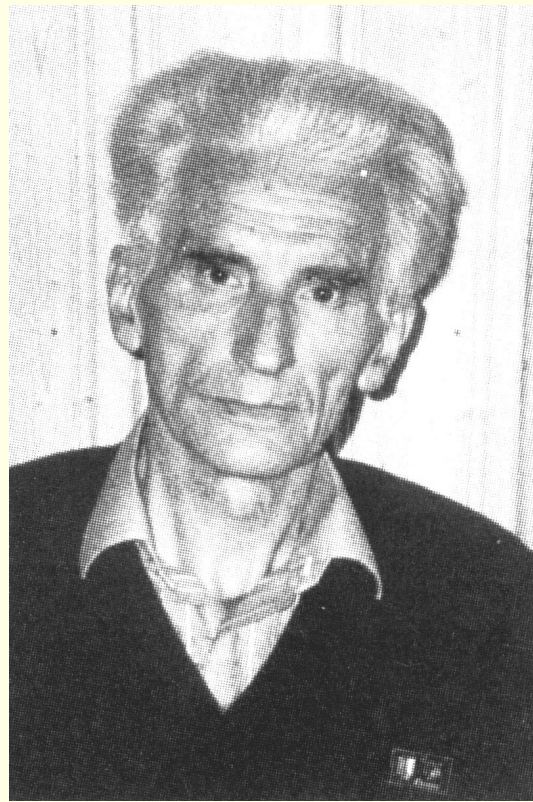
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Evolution of the Project:

- **Plan A: Integrate the Primitive Equations**
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- **Plan B: Integrate baroclinic Q-G System**
Too computationally demanding
- **Plan C: Solve barotropic vorticity equation**
Very satisfactory initial results

Charney, Fjørtoft, von Neumann



Charney, et al., *Tellus*, 1950.

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$$\left[\begin{array}{c} \text{Absolute} \\ \text{Vorticity} \end{array} \right] = \left[\begin{array}{c} \text{Relative} \\ \text{Vorticity} \end{array} \right] + \left[\begin{array}{c} \text{Planetary} \\ \text{Vorticity} \end{array} \right] \quad \eta = \zeta + f.$$

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The atmosphere is treated as a single layer, and the flow is assumed to be nondivergent. **Absolute vorticity is conserved:**

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In more detail:

$$\frac{\partial}{\partial t}[\nabla^2 \psi - F\psi] + \left\{ \frac{\partial \psi}{\partial x} \frac{\partial \nabla^2 \psi}{\partial y} - \frac{\partial \psi}{\partial y} \frac{\partial \nabla^2 \psi}{\partial x} \right\} + \beta \frac{\partial \psi}{\partial x} = 0$$

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- Timestep : $\Delta t = 1$ hour (2 and 3 hours also tried)
 - Gridstep : $\Delta x = 750$ km (approximately)
 - Gridsize : $18 \times 15 = 270$ points
 - Elapsed time for 24 hour forecast: About 24 hours.

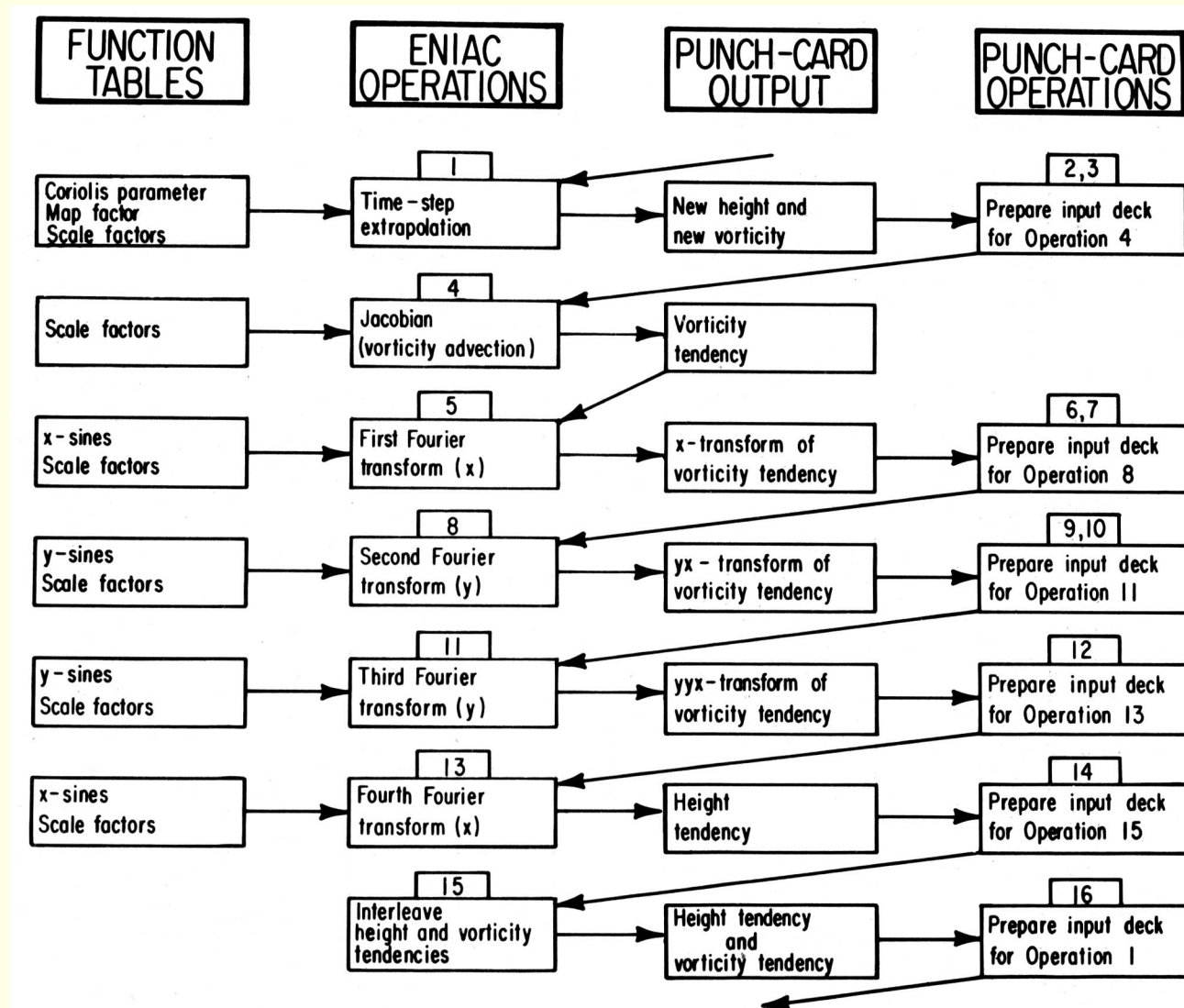
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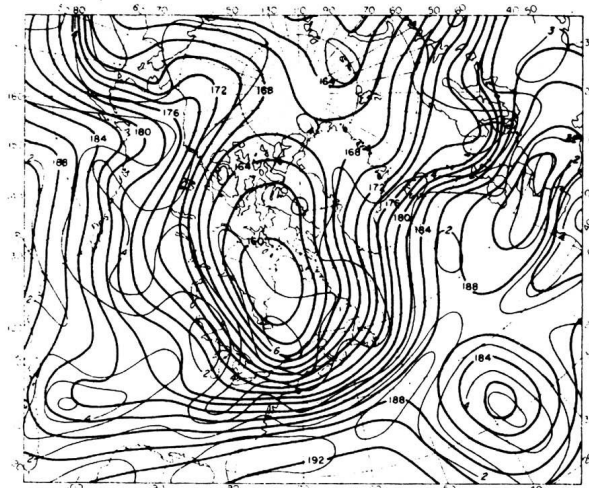
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Forecast involved **punching about 25,000 cards**. Most of the elapsed time was spent handling these.

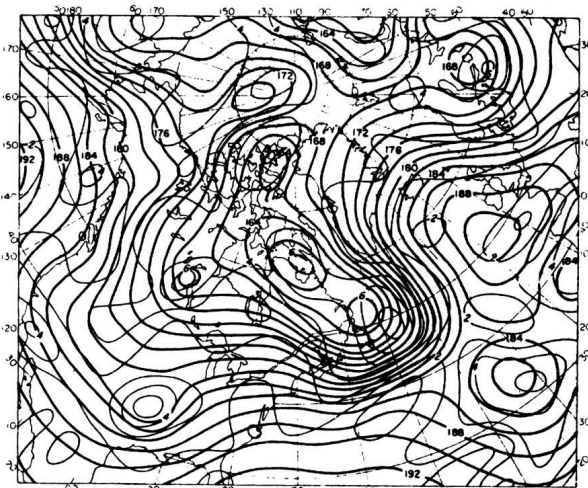
ENIAC Algorithm



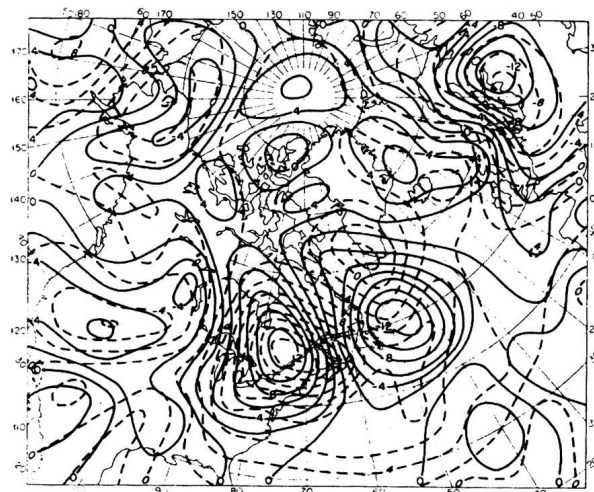
ENIAC: First Computer Forecast



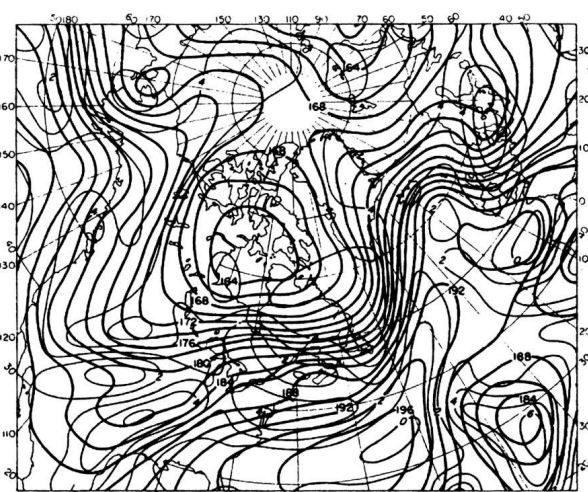
(A)



(B)



(C)



(D)

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- “This is . . . **an enormous scientific advance** on the single, and quite wrong, result in which . . . [Richardson (1922)] ended.”