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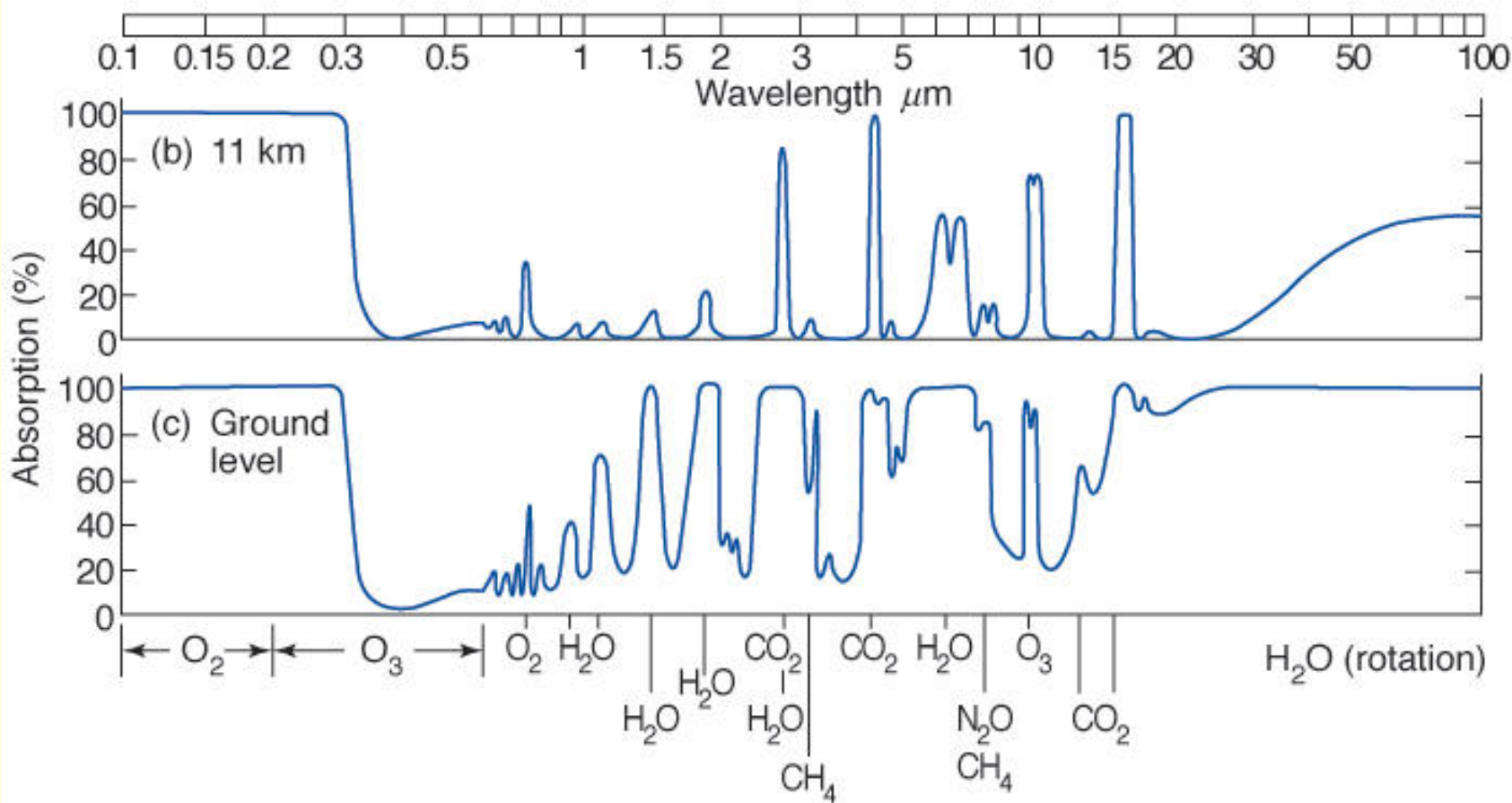
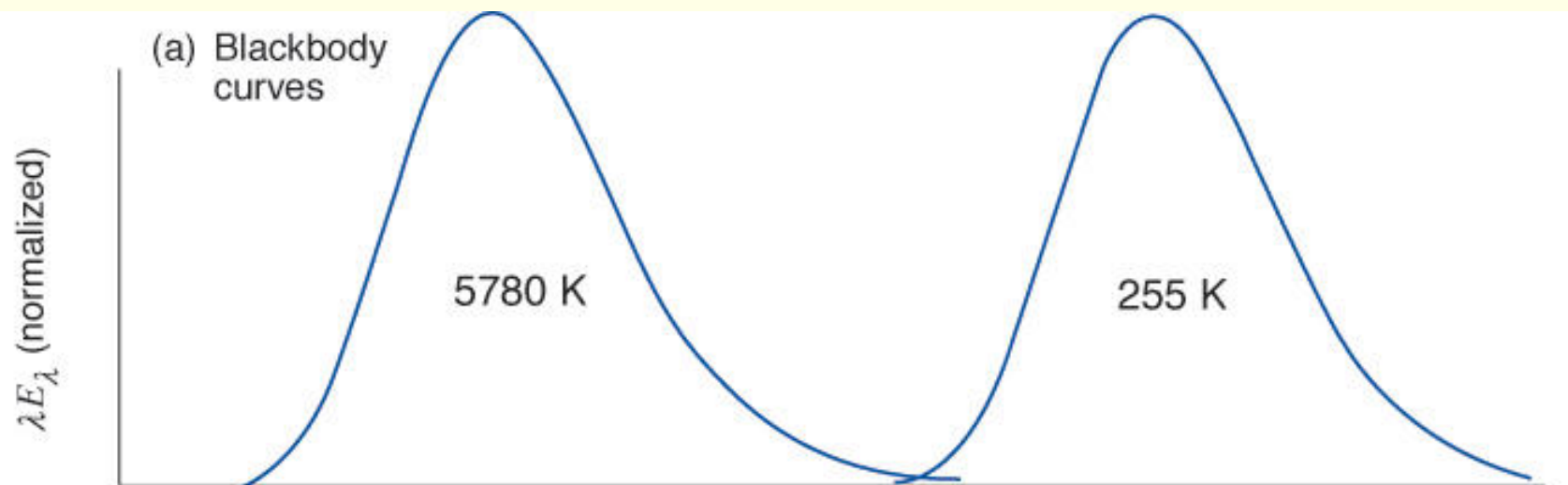
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The distinction is quite striking, as shown in the following figure.



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Explain why the surface temperature computed above is considerably higher than the effective temperature in the absence of an atmosphere.

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This must equal the upward flux from the surface:

$$F_E = 0.9 \times F_S + F_L$$

The upward and downward fluxes at the top of the atmosphere must also be in balance, which gives us the relation

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$$\begin{aligned} F_L - F_E &= -0.9F_S \\ F_L + 0.2F_E &= F_S \end{aligned}$$

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$$\sigma T_{\text{surface}}^4 = F_E = 380 \text{ W m}^2$$

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$$\sigma T_{\text{surface}}^4 = F_E = 380 \text{ W m}^2$$

Therefore, since $\sigma = 5.67 \times 10^{-8} \text{ W m}^{-2} \text{ K}^{-4}$, we have

$$T_{\text{surface}} = \sqrt[4]{\frac{380}{5.67 \times 10^{-8}}} = 286 \text{ K} = +13^\circ \text{C}$$

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Note that the surface temperature in this case is some 31°C higher than in the case of exercise 4.6 when there was no atmosphere:

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No atmosphere:

$$T_{\text{surface}} = -18^\circ\text{C}$$

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Show how the surface temperature of the planet is affected by the presence of this atmosphere and describe the radiative equilibrium temperature profile in the atmosphere of the planet.

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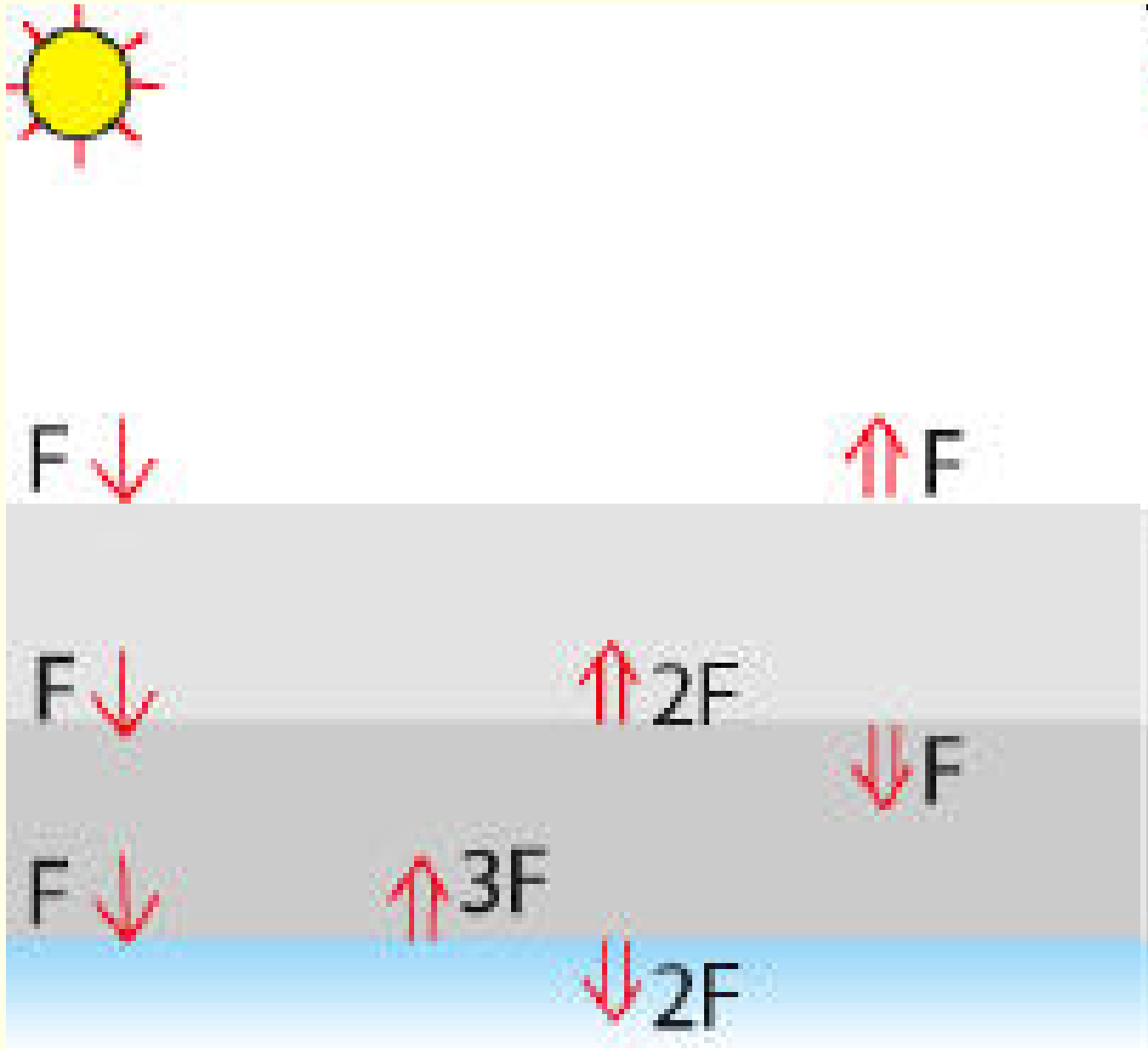
To balance the incident radiation, the lower layer must emit $2F$ units of longwave radiation. Since the layer is isothermal, it also emits $2F$ units of radiation in the downward radiation.

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To balance the incident radiation, the lower layer must emit $2F$ units of longwave radiation. Since the layer is isothermal, it also emits $2F$ units of radiation in the downward radiation.

Hence, the downward radiation at the surface of the planet is F units of incident solar radiation plus $2F$ units of longwave radiation emitted from the atmosphere, a total of $3F$ units, which must be balanced by an upward emission of $3F$ units of longwave radiation from the surface.



Radiation balance for a planetary atmosphere that is transparent to solar radiation and consists of two isothermal layers that are opaque to planetary radiation.

By induction, the above reasoning can be extended to an N -layer atmosphere.

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The emissions from the atmospheric layers, working downward from the top, are $F; 2F; 3F:::NF$ and the corresponding radiative equilibrium temperatures are 255, 303, 335....
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To estimate the corresponding radiative equilibrium lapse rate within the atmosphere we would need to take into account the fact that the geometric thickness of opaque layers decreases rapidly as one descends through the atmosphere owing to the increasing density of the absorbing media with depth.

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Hence, the radiative equilibrium lapse rate steepens with increasing depth.

In effect, radiative transfer becomes less and less efficient at removing the energy absorbed at the surface of the planet due to the increasing blocking effect of the greenhouse gases.

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The bottom panel of Fig. 4.5 shows that the wavelength dependence is quite pronounced, with well defined *absorption bands* identified with specific gaseous constituents, interspersed with *windows* in which the atmosphere is relatively transparent.

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- The water evaporated from the ocean is carried upward by **convection**.
- The moisture reaches levels **above** the main infra-red absorbers.
- The **latent heat** is then released by condensation, from where much of it radiates to space.

End of §4.3